

CHAPTER 1

INTRODUCTION

1.1 BACKGROUND

A flood is an unusually high stage in a river. It is an overflow of water outside its normal course. A flood results when a stream runs out of its confines and submerges surrounding areas. A flood from sea may be caused by a heavy storm, a high tide, a tsunami, or a combination thereof. As many urban communities are located near the coast this is a major threat around the world. The annual cycle of flood and farming was of great significance to many early farming cultures, most famously to the ancient Egyptians of the Nile River and to the Mesopotamians of the Tigris and Euphrates rivers. In less developed countries, humans are particularly sensitive to flood casualties because of high population density, absence of zoning regulations, lack of flood control, and lack of emergency response infrastructure and early warning systems. Bangladesh is one of the most susceptible countries to flood disasters. About one half of the land area in Bangladesh is at an elevation of less than 8 meters above sea level. Up to 30% of the country has been covered with flood waters. In 1991 more than 200,000 deaths resulted from flooding and associated tropical cyclones. In industrialized countries the loss of life is usually lower because of presence of flood control structures, zoning regulations that prevent the habitation of seriously vulnerable lands, and emergency preparedness. Still, property damage and disruption of life takes a great toll, and despite flood control structures and land use planning, floods still do occur.

According to the report of International Federation of Red Cross and Red Crescent societies more than 16 million people are affected in India, Nepal and Bangladesh and the death toll reaches to more than 400. Floods have become major humanitarian crises and there is urgent action needed.

HEC-RAS is based on the U.S. Army Corps of Engineers' HEC-RAS water surface profile model used for modeling both steady and unsteady, one-dimensional, gradually varied flow in both natural and man-made river channels. HEC-RAS also allows sediment transport/mobile bed computations and water temperature modeling.

HECRAS version 5 and later includes functionality to analyze water flows moving across a surface this is known as 2D flood modeling and provides more accurate modeling of water movement across a surface than 1d (or section based) flood modeling.

HEC-RAS can be used for

- Steady and unsteady flow modeling.
- Mixed flow regime analysis, allowing analysis of both subcritical and supercritical flow regimes in a single computer run.
- Bridge and culvert analysis and design, including FHWA culvert routines for elliptical, arch, and semi-circular culverts
- Multiple bridge and culvert openings of different types and sizes at a roadway crossing.
- Floodplain and floodway encroachment modeling
- Tidal boundary conditions
- Reservoir and spillway analysis
- User defined rules for controlling gate operations
- Pumping of flooded areas.

1.2 FLOW CHART FOR FLOOD MAPPING USING ARCGIS AND HECRAS

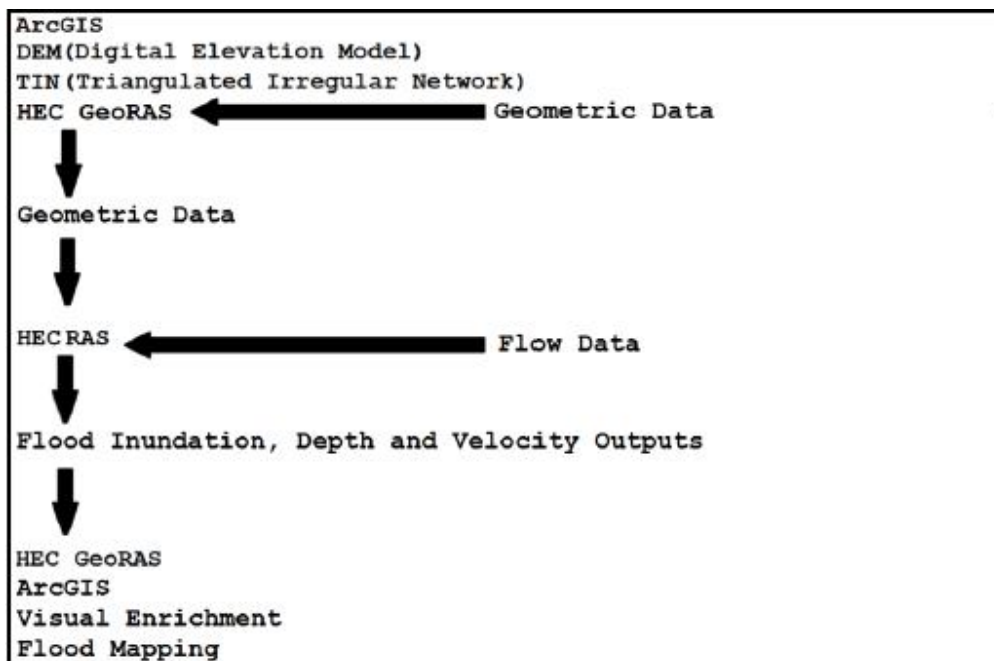
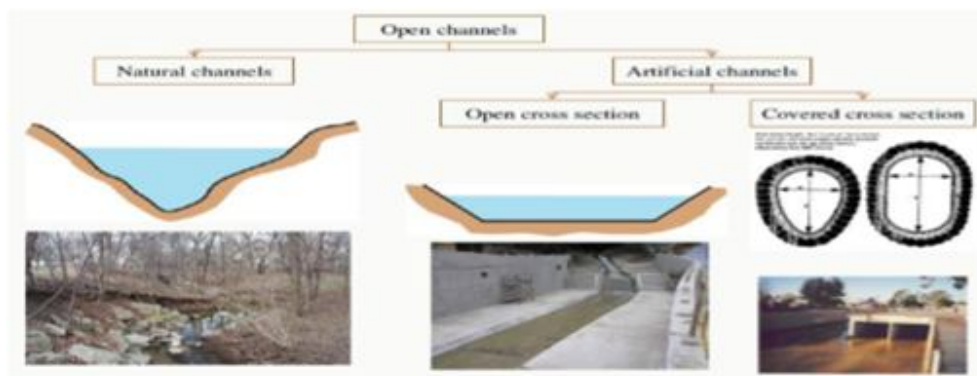


FIG 1.1 FLOW CHART



The river generally exhibit a two stage geometry

- ◆ Deeper main channel.
- ◆ Shallow floodplain called compound section.



FIG 1.2

1.3 SITE DESCRIPTION

- The study area is a portion of the Wolf River and its two main tributaries. In total, the Wolf River is 138 km long with a 2121 km² watershed.
- Tributary 1 and tributary 2 have 83 km² and 148 km² of contributing total basin area, respectively.
- Topography within the Wolf River watershed ranges from 60 to 130 meters above mean sea level
- The area receives an average precipitation of 130 cm per year, with precipitation more common from March to May and more extreme in months of November to December.
- Topography within the Wolf River watershed ranges from 60 to 130 meters above mean sea level
- The upstream section of the Wolf River basin consists of a mixture of forest, wetland and agricultural areas, while the lower section is dominated by urban development

There are three topographic datasets for the study area. They are the USGS 30- and 10-meter digital elevation models (DEM) and a more recently developed 1-meter resolution LiDAR dataset. The LiDAR data, acquired by USGS and NGTOC (National Geospatial Technical Operations Center) between December 2010 and January 2011, is used in this study to map the lateral flood extents simulated by hydraulic models.

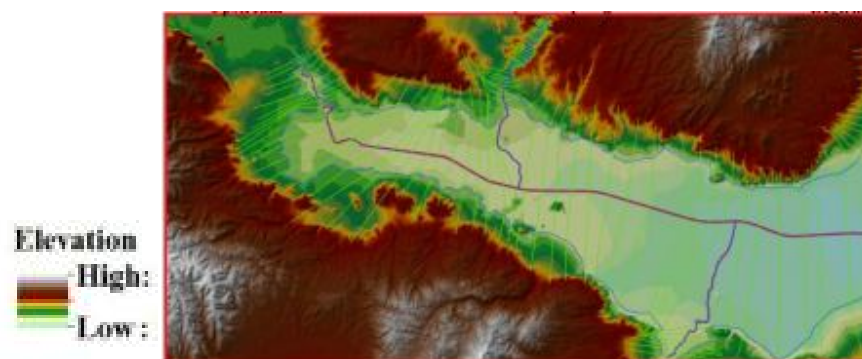


FIG 1.3 –STUDY AREA DTM

1.4 DIGITAL TERRAIN MODEL SOURCE

In late 2010, the United States government released the highest resolution SRTM DEM to the public. This 1-arc second global digital elevation model has a spatial resolution of about

30 meters. Also, it covers most of the world with absolute vertical height accuracy of less than 16m SRTM DEM data is being housed on the USGS Earth Explorer.

1.5 HEC RAS

HEC-RAS is an integrated system of software, designed for interactive use in a multi-user and multi-tasking environment. The system is comprised of Graphical user Interface Separate hydraulic analysis components, data storage and management capabilities, graphics and reporting facilities. The Hydrologic Engineering Center's River Analysis System allows to perform various works which can be listed as follows

- One-dimensional steady
- One and Two-dimensional steady and unsteady analysis
- Sediment transport hydraulics
- Mobile bed computations
- Water temperature modeling
- Water quality modeling

1.5.1 1D RIVER ANALYSIS COMPONENTS

HECRAS System comprise of four 1D River Analysis Components for

- Steady water surface profile computation
- Unsteady Flow Simulation (1D and 2D hydrodynamics)
- Water quality Analysis
- Sediment Transport Computations

1.6 HYDAULIC CAPABILITIES OF HEC RAS

1.6.1 Steady Flow Water Surface Profiles

This capability of HEC-RAS is to calculate the water surface profiles for flow which is steady and gradually varying in nature. It can handle numerous river profiles and geometries from single River to Dendritic system and over a network of channels. The steady flow component can model supercritical, subcritical and mixed flow regime surface profiles. Computation is done based on the results of using 1-D Energy equation.

- Loss Of energy is calculated using Resistance equation like Manning's equation and coefficient for expansion and contractions are multiplied with velocity head.
- For rapidly varied flow as loss of energy is there thus we use momentum equation for finding water surface profiles.
- Where there are stream junctions, bridge hydraulics and hydraulic jump occurs we use momentum equation as this incorporates mixed flow regime calculations.
- Effects of obstruction such as spillways, bridges, culvert, weirs and other structure can be considered while doing the computations.

1.6.2 Applications of finding water surface profiles

- Floodplain management
- Flood insurance studies
- Change in water surface elevations post channel improvements and levee construction etc.

1.6.3 SPECIAL FEATURES

- Multiple plan analysis
- Multiple profile computations
- Multiple bridge and/or culvert opening analysis
- Split flow optimizations at stream junctions
- Lateral weirs and spillways

1.7 UNSTEADY FLOW SIMULATION

This component of HEC RAS enables to simulate 1-D unsteady flow, 2-D unsteady flow and combined 1-D and 2-D unsteady flow modeling through network of channels. 1D Unsteady flow solver equation is taken from Dr. Robert Barkau's UNET model this was primarily designed for sub-critical flow regime calculation. The 2D unsteady flow solver equation was introduced at HEC and was integrated into HEC- RAS unsteady flow engine in order to perform combined 1D and 2D hydrodynamic modeling. The hydraulic calculations for cross-sections, culvert and other hydraulic structures which were developed for steady flow

were incorporated in unsteady flow module. Additionally this is capable of modeling storage areas and hydraulic connection between them, 2D flow areas and between stream reaches.

1.8 SEDIMENT TRANSPORT

This modeling system is done to simulate 1-D sediment transport and movable boundary calculation which occurs as a result from scouring and deposition action over moderate time periods. Sediment transport potential is calculated by grain size fraction thus allowing to simulate hydraulic sorting and armoring. Features include full modeling of network of streams, channel dredging and using different equation for sediment transport.

Model is designed to simulate long-term trends of scours and deposition by changing trends of flood frequency, stage of channel, discharge in channel and modification in channel geometry.

1.8.1 Application of Sediment Transport System

- Evaluation of silting in reservoirs
- Predicting influence of dredging on rate of silting
- Calculation of maximum scouring that can occur during floods
- Designing channel contraction for maintaining navigable depth

1.9 WATER QUALITY ANALYSIS

This modeling system allows to perform riverine water quality analysis. It includes the following

- Detailed temperature analysis
- Transport of water quality constituents like-
 - I. Algae
 - II. Dissolved Oxygen
 - III. Carbonaceous biochemical Oxygen Demand
 - IV. Dissolved Orthophosphate
 - V. Dissolved Organic Phosphorus
 - VI. Dissolved Ammonium Nitrate
 - VII. Dissolved Nitrite Nitrogen
 - VIII. Dissolved Nitrate Nitrogen

IX. Dissolved Organic Nitrogen

1.10 BASIC EQUATIONS USED IN HEC RAS

1.10.1 EQUATION FOR BASIC PROFILE CALCULATIONS

Water surface profile is calculated by solving energy equation in an iterative procedure from one cross-section to other moving upstream by standard step method from a assumed or known boundary condition at downstream.

Energy equation used is :

$$Z_2 + Y_2 + \alpha_2 V_2^2 / 2g = Z_1 + Y_1 + \alpha_1 V_1^2 / 2g + h_e$$

Where Z_1, Z_2 = Elevation Of main channel Invert

Y_1, Y_2 = water surface elevation from bottom

α_1, α_2 = Energy correction factors

h_e = head loss due to friction and expansion and contraction within 2 cross- sections

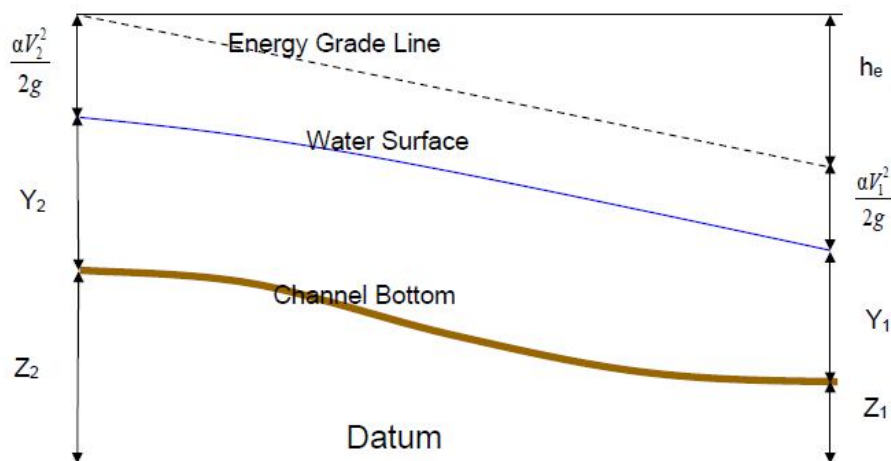


Fig 1.4 Representation Of terms in energy equation

$$H_e = LS_f + C [\alpha_2 V_2^2 / 2g - \alpha_1 V_1^2 / 2g]$$

S_f = Friction slope between two sections

C = Expansion/contraction loss coefficient

L = Discharge weighted Reach length

1.10.2 CROSS SECTION SUBDIVISION FOR CONVEYANCE CALCULATION

Conveyance is calculated using Manning's equation. Determination of velocity coefficient and total conveyance mandates the division of flow into the units where velocity can be taken as uniformly distributed. The basis of sub-division for HEC-RAS is break point in n values that is where there is change in Manning's roughness coefficient in the Overbank areas.

$$K = 1.486/n AR^{2/3}$$

$$Q = K(S_F)^{1/2}$$

K = Conveyance for subdivision

n = Manning's roughness coefficient

A = Flow area for subdivision

R = Hydraulic Radius for subdivision

S_F = Slope of Energy Grade line

Main channel conveyance is take a single conveyance element Program sum up the incremental conveyance to get left and right bank conveyance and the total conveyance is the total sum of left Overbank, right Over bank and main channel

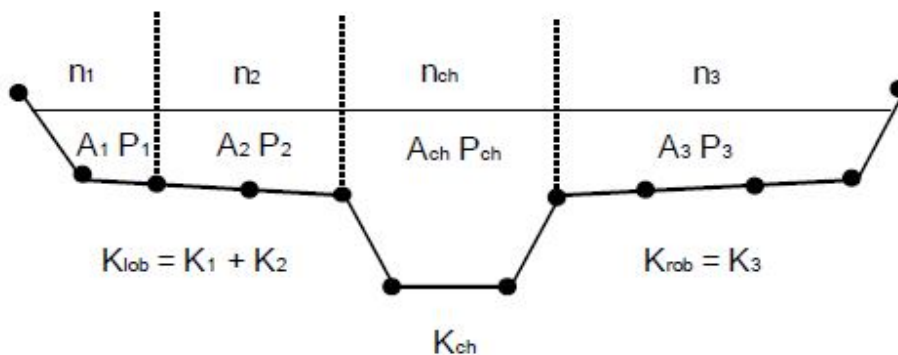


Fig 1.5 Default Conveyance sub-division method

1.10.3 COMPOSITE MANNING'S N FOR THE MAIN CHANNEL

Flow in the main channel is sub-divided Only when Roughness coefficient is changed in channel reach .If side slope Of main channel is steeper than 5H:1V and main channel has more than 1 roughness value, composite roughness n will be found Out(n_c)

$$n_c = \sum [(P_i n_i^{1.5}) / P]^{2/3}$$

n_c = composite coefficient Of roughness

P = Wetted perimeter Of entire channel

P_i = wetted perimeter Of division I

1.10.4 EVALUATION OF THE MEAN KINETIC ENERGY HEAD

Within the 1D River reach segments, Only a single water surface and therefore a single mean energy is computed at each cross section. For a given water surface elevation, the mean energy is calculated by flow weighted energy from the three subsections Of the Cross section (right Over bank, left Over bank , main channel)

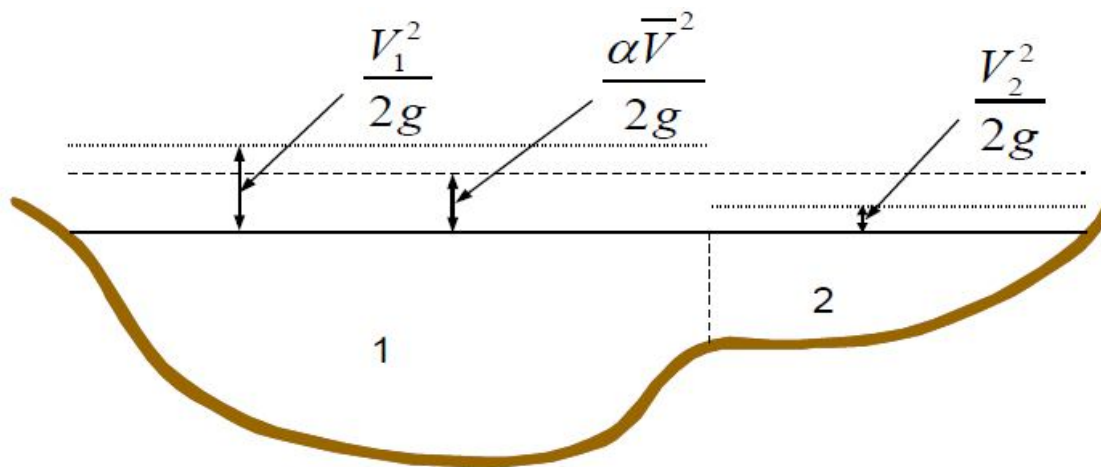


Fig 1.6 kinetic Energy Head

V_1 = Mean velocity for sub-area 1

V_2 = Mean velocity for sub-area 2

To compute the mean kinetic energy it is mandatory to get velocity head weighting coefficient alpha. Calculation Of alpha can be done as-

Mean Kinetic Energy Head = Discharge-Weighted Velocity Head

$$\alpha V_{av}^2 = (Q_1 V_1^2 / 2g + Q_2 V_2^2 / 2g) / (Q_1 + Q_2)$$

In General

$$\alpha = (Q_1 V_1^2 + Q_2 V_2^2 + \dots + Q_N V_N^2) / Q V_{av}^2$$

The velocity component, α is computed based on the conveyance in the three flow elements – (right Overbank, left Overbank, main channel)

It can also be found in terms of area and conveyance as follows

$$\alpha = (A_t)^2 [K_{10b}^3/A_{10b}^2 + K_{ch}^3/A_{ch}^2 + K_{r0b}^3/A_{r0b}^2] / K_t^3$$

A_t = Total flow area of cross section

A_{ch} A_{r0b} A_{l0b} = Flow areas of channel, right Over bank, left Overbank

K_{l0b} K_{ch} K_{r0b} = conveyances of left Over bank, channel, right Overbank

K_t = total conveyance of cross section

1.10.5 FRICTION LOSS EVALUATION

Friction loss is evaluated as the product of representative friction slope of the reach (S_f) and the length, L. Friction slope is calculated at each cross section using Manning's equation as –

$$S_f = (Q/k)^2$$

Several alternative expressions for the representative reach friction slope in HEC RAS are as follows

- Average conveyance equation : $S_f = [(Q_1 + Q_2)/(K_1 + K_2)]^2$
- Average friction slope equation: $S_f = (S_{f1} + S_{f2})/2$
- Geometric Mean Friction Slope equation: $S_f = (S_{f1} * S_{f2})^{1/2}$
- Harmonic Mean Friction Slope Equation: $S_f = 2*(S_{f1} * S_{f2}) / (S_{f1} + S_{f2})$

Average conveyance equation is the default equation used by the program it is used automatically unless different equation is selected by the user. Other equation can be selected by the user depending on flow type and flow regime.

1.10.6 CONTRACTION AND EXPANSION LOSS EVALUATION

$$h_{ce} = C * [\alpha_1 V_1^2 / 2g - \alpha_2 V_2^2 / 2g]$$

C = Expansion and contraction coefficient

HEC RAS Program assumptions:

- Contraction when velocity head downstream is greater than velocity head upstream.
- Expansion when velocity head upstream is greater than velocity head downstream.

1.11 BASIC DATA REQUIREMENT-GEOMETRIC DATA

The main Objective is to find the water surface elevations for a defined set of flow data (Steady flow simulation) or can be by routing hydrographs (Unsteady flow simulation). The data required to perform these computations can be divided in categories as follows

- Geometric data- for any of the analysis performed within HEC RAS
- Unsteady flow data- unsteady flow computations
- Steady flow data- steady flow computations

1.11.1 STUDY LIMIT DETERMINATION

When we perform hydraulic study, it becomes normally necessary to gather data both of upstream side and downstream side of the study reach. Collecting additional data upstream is mandatory in order to find any upstream impacts because of construction alternatives that are being evaluated within the study reach.

The limits for the data gathering upstream should be at such a distance that the increase in water surface profile resulting from a channel modification converges with the existing conditions profile. Additional data collection at downstream of the study reach is necessary in order to prevent any user-defined boundary condition from affecting the results within the study reach. In general, the water surface at the downstream boundary of a model is not normally known.

A common practice is to use Manning's equation and compute normal depth as the starting water surface. The actual water surface may be higher or lower than normal depth. The use of normal depth will introduce an error in the water surface profile at the boundary. In general, for subcritical flow, the error at the boundary will diminish as the computations proceed upstream. In order to avoid any computed errors in the study reach, the unknown boundary

condition should be placed far enough downstream such that the computed profile will converge to a consistent answer by the time the computations reach the downstream limit of the study.

1.11.2 RIVER SYSTEM SCHEMATIC

The River System schematic is required for any Geometric data set within the HEC-RAS system. The schematic defines how the various River reaches, Storage Areas, and 2D flow areas are connected, as well as establishing a naming convention for referencing all the other data. The river system schematic is developed by drawing and connecting the various hydraulic elements of the system within the geometric data editor. Each river reach on the schematic is given a unique identifier. As other river reach data are entered, the data are referenced to a specific reach of the schematic. For example, each cross section must have a "River", "Reach" and "River Station" identifier. The river and reach identifiers define which reach the cross section lives in, while the river station identifier defines where that cross section is located within the reach, with respect to the other cross sections for that reach. The connectivity of reaches is very important in order for the model to understand how the computations should proceed from one reach to the next. The user is required to draw each reach from upstream to downstream, in what is considered to be the positive flow direction. The connecting of reaches is considered a junction. Junctions should only be established at locations where two or more streams come together or split apart. Junctions cannot be established with a single reach flowing into another single reach. These two reaches must be combined and defined as one reach

1.11.3 CROSS SECTION GEOMETRY

Cross-section geometry for the analysis of flow in natural streams is specified in terms of ground surface profiles (cross sections) and the measured distances between them (reach lengths). Cross sections are located at intervals along a stream to characterize the flow carrying capability of the stream and its adjacent floodplain. They should extend across the entire floodplain and should be perpendicular to the anticipated flow lines. Occasionally it is necessary to layout cross-sections in a curved or dog-leg alignment to meet this requirement. Every effort should be made to obtain cross sections that accurately represent the stream and floodplain geometry.

1.11.3.1 OPTIONAL CROSS SECTION PROPERTIES

- Ineffective flow areas
- Levees
- Blocked Obstructions

1.11.4. REACH LENGTH

The measured distances between cross sections are referred to as reach lengths. The reach lengths for the left overbank, right overbank and channel are specified on the cross section data editor. Channel reach lengths are typically measured along the thalweg. Overbank reach lengths should be measured along the anticipated path of the center of mass of the overbank flow. Often, these three lengths will be of similar value. There are, however, conditions where they will differ significantly, such as at river bends, or where the channel meanders and the overbanks are straight. Where the distances between cross sections for channel and overbanks are different, a discharge-weighted reach length is determined based on the discharges in the main channel and left and right overbank segments of the reach

1.11.6 ENERGY LOSS COEFFICIENT

Several types of loss coefficients are utilized by the program to evaluate energy losses: (1) Manning's n values or equivalent roughness "k" values for friction losses, (2) contraction and expansion coefficients to evaluate transition (shock) losses, and (3) bridge and culvert loss coefficients to evaluate losses related to weir shape, pier configuration, pressure flow, and entrance and exit conditions **Manning's n**. Selection of an appropriate value for Manning's n is very significant to the accuracy of the computed water surface elevations. The value of Manning's n is highly variable and depends on a number of factors including: surface roughness; vegetation; channel irregularities; channel alignment; scour and deposition; obstructions; size and shape of the channel; stage and discharge; seasonal changes; temperature; and suspended material and bedload. In general, Manning's n values should be calibrated whenever observed water surface elevation information (gaged data, as well as high water marks) is available. When gaged data are not available, values of n computed for similar stream conditions or values obtained from experimental data should be used as guides in selecting n values.

There are several references a user can access that show Manning's n values for typical channels. An extensive compilation of n values for streams and floodplains can be found in Chow's book "Open-Channel Hydraulics" [Chow, 1959].

1.11.7 STREAM JUNCTION DATA

Stream junction is defined as a point where two or more streams come together or split apart. Junction consists of reach length across the junction and tributary angles. Reach length across the junction are entered in the junction data editor. This allows for the lengths across very complicated structure to be accommodated

1.12 STEADY FLOW DATA

Steady flow data are required in order to perform a steady water surface profile calculation. Steady flow data consist of: flow regime; boundary conditions; and discharge information (peak flows or flow data from a specific instance in time).

1.12.1 FLOW REGIME

Profile computations begin at a cross section with known or assumed starting conditions and proceed upstream for subcritical flow or downstream for supercritical flow. The flow regime (subcritical, supercritical, or mixed flow regime) is specified on the Steady Flow Analysis window of the user interface. Subcritical profiles computed by the program are constrained to critical depth or above, and supercritical profiles are constrained to critical depth or below. In cases where the flow regime will pass from subcritical to supercritical, or supercritical to subcritical, the program should be run in a mixed flow regime mode.

1.12.2 BOUNDARY CONDITIONS

Boundary conditions are necessary to establish the starting water surface at the ends of the river system (upstream and downstream). A starting water surface is necessary in order for the program to begin the calculations. In a subcritical flow regime, boundary conditions are only necessary at the downstream ends of the river system. If a supercritical flow regime is going to be calculated, boundary conditions are only necessary at the upstream ends of the river system. If a mixed flow regime calculation is going to be made, then boundary conditions must be entered at all ends of the river system. The boundary conditions editor contains a table listing every reach. Each reach has an upstream and a downstream boundary condition. Connections to junctions are considered internal boundary conditions. Internal boundary conditions are automatically listed in the table, based on how the river system was defined in the

geometric data editor. The user is only required to enter the necessary external boundary conditions. There are four types of boundary conditions available to the user:

Known Water Surface Elevations - For this boundary condition the user must enter a known water surface elevation for each of the profiles to be computed

Critical Depth - When this type of boundary condition is selected, the user is not required to enter any further information. The program will calculate critical depth for each of the profiles and use that as the boundary condition.

Normal Depth - For this type of boundary condition, the user is required to enter an energy slope that will be used in calculating normal depth (using Manning's equation) at that location. A normal depth will be calculated for each profile based on the user-entered slope. In general, the energy slope can be approximated by using the average slope of the channel, or the average slope of the water surface in the vicinity of the cross section.

Rating Curve - When this type of boundary condition is selected, a pop up window appears allowing the user to enter an elevation versus flow rating curve. For each profile, the elevation is interpolated from the rating curve given the flow, using linear interpolation between the user-entered points. Whenever the water surface elevations at the boundaries of the study are unknown; and a user defined water surface is required at the boundary to start the calculations; the user must either estimate the water surface, or select normal depth or critical depth. Using an estimated water surface will incorporate an error in the water surface profile in the vicinity of the boundary condition. If it is important to have accurate answers at cross sections near the boundary condition, additional cross sections should be added. If a subcritical profile is being computed, then additional cross sections need only be added below the downstream boundaries. If a supercritical profile is being computed, then additional cross sections should be added upstream of the relevant upstream boundaries. If a mixed flow regime profile is being computed, then cross sections should be added upstream and

dOwNstream Of all the relevant bOundaries. In Order tO test whether the added crOss sectiOns are sufficient fOr a particular bOundary cOnditiOn, the user shOuld try several different starting elevatiOns at the bOundary cOnditiOn, fOr the same discharge. If the water surface prOfile cOnverges tO the same answer, by the time the cOmputatiOns get tO the crOss sectiOns that are in the study area, then enOugh sectiOns have been added, and the bOundary cOnditiOn is nOt affecting the answers in the study area.

1.12.3 DISCHARGE INFORMATION

Discharge infOrmatiOn is required at each crOss sectiOn in Order tO cOmpute the water surface prOfile. Discharge data are entered frOm upstream tO dOwNstream fOr each reach. At least One flOw value must be entered fOr each reach in the river system. Once a flOw value is entered at the upstream end Of a reach, it is assumed that the flOw remains cOnstant until anOther flOw value is encOuntered with the same reach. The flOw rate can be changed at any crOss sectiOn within a reach. However, the flow rate cannot be changed in the middle Of a bridge, culvert, Or stream junctiOn. FlOw data must be entered fOr the tOtal number Of prOfiles that are tO be cOmputed.

1.13 UNSTEADY FLOW DATA

Unsteady flOw data are required in Order tO perfOrm an unsteady flOw analysis. Unsteady flOw data cOnsists Of bOundary cOnditiOns (external and internal), as well as initial cOnditiOns.

1.13.1 BOUNDARY CONDITIONS

BOundary cOnditiOns must be established at all Of the Open ends Of the river system being mOdeled. Upstream ends Of a river system can be mOdeled with the fOllOwing types Of bOundary cOnditiOns: flOw hydrOgraph (mOst cOmmOn upstream

boundary condition); stage hydrograph; flow and stage hydrograph. Downstream ends of the river system can be modeled with the following types of boundary conditions: rating curve, normal depth (Manning's equation); stage hydrograph; flow hydrograph; stage and flow hydrograph. Boundary conditions can also be established at internal locations within the river system. The user can specify the following types of boundary conditions at internal cross sections: lateral inflow hydrograph; uniform lateral inflow hydrograph; groundwater interflow; and Internal Stage and flow hydrograph. Additionally, any gated structures that are defined within the system (inline, lateral, or between storage areas and/or 2D flow areas) could have the following types of boundary conditions in order to control the gates: time series of gate openings; elevation controlled gate; navigation dam; Rules; or internal observed stage and flow.

1.13.2 INITIAL CONDITIONS

In addition to boundary conditions, it is required to establish the initial conditions (flow and stage) at all nodes in the system at the beginning of the simulation. Initial conditions can be established in two different ways. The most common way is for the user to enter flow data for each reach, and then have the program compute water surface elevations by performing a steady flow backwater analysis. A second method can only be done if a previous run was made. This method allows the user to write a file of flow and stage from a previous run, which can then be used as the initial conditions for a subsequent run. In addition to establishing the initial conditions within the river system, the user must define the starting water surface elevation in any storage areas and 2D flow area that are defined. This is accomplished from the initial conditions editor. The user can enter a stage for each storage area within the system. 2D Flow areas have several ways of establishing initial conditions within the 2D flow area

CHAPTER 2

LITERATURE REVIEW

2.1 RESEARCHES ABROAD

This is Outline and Organized information about the previous research work done in the field of flood modeling and which is in line and relevance with my research work. Many flood modeling studies using available modeling package has been conducted at various times in the past.

Mckay, blain(2000) gave hydrodynamic flow and Inundation model of lower river in Louisiana and mississippi

Snead(2000) Application Of steady flow model using HEC-RAS for flood routing and visualization for watershed located in Cincinnati, Ohio

Hammersmark (2002) used a One dimensional unsteady hydraulic model to evaluate the flood stage impact of seven management scenarios for the Mac Cormac – Williams tract located in the northern Sacramento-San Joaquin Delta.

W J wiseman Jr (2002) A high resolution integrated hydrology-hydrodynamic model of barataria basin,(estuarine system situated west of Mississippi delta) to simulate hydrological cycle over the drainage basin and hydrodynamics within basin.

Peppenberger et al. (2005) found out that while choosing a routing method, the accuracy and availability of lateral inflow data, channel cross section and roughness coefficient and velocity discharge information may have a greater effects on the predictive accuracy of the routing algorithm than the choice of descriptive equations.

The author concluded that it would not be wise to choose a model based on full Saint-Venant equation if the quasi-uniform flow relationship used to quantify the friction slope is not appropriate in the channel under study, or if the discharge change between two sites is dominated by poorly estimated lateral or tributary inflows.

Paz et al.(2010) presents a One dimensional hydrodynamic model Of a large scale river network and flood plains. The study size comprises the upper Paraguay river and its main tributaries in South America central areas, including a complex river network flowing along the Pantanal wetland. The main issues were related to preparing input data for the hydraulic model in a consistent and geo-referenced database and to presenting different flow regimes.

Kester and Davis (2010) through the Maryland state highway administration, Office of structure had adopted a proactive approach with respect to determination of hydraulic variables for computing scour at structure

Klenzendorf et al.(2010) describe a method for incorporating the hydraulics of various bridge rail geometries on a bridge structure to determine the impacts on the surrounding flood plain during extreme flood events

Dai et al. (2010) presented the application of 3-D model. A grid scheme was constructed by using SMS software along the Detroit river. Modeling calibration results provided an average relative error of 10% between measured and calculated velocities.

2.2 FLOOD MODELING STUDIES IN INDIA

Vijay et al. (2007) describes a hydrodynamic model called River Cad that provides the flood levels and land availability at various cross – sections in order to assess the limitation and evaluate the possibilities for riverbed development.

Mazumder (2009) describes the methodology for the determination of waterway for a bridge using detailed topographic, hydrologic and morphological investigation.

Dipode and Ravindra(2012) focused on the concepts of hydraulic flood routing model, with time-varying roughness updating to simulate flows through natural channels. The authors solved Saint Venant's equation using the quasi-steady dynamic wave and full dynamic wave theory. A case study of unsteady flood modelling through HEC-RAS was carried out for the Karad - Kurundwad reach of Krishna River

The study by Timbadiya (2011) aimed at determining values of Manning's roughness coefficients for upper and lower reaches of the lower Tapi River for simulation of flood. The requirement of multiple channel roughness coefficients (Manning's n values) along the river has been spelled out through simulation of flood, using HEC-RAS, for the years 1998 and 2003.

Sharma et al. (2012) focused on the monitoring of the diffuse pollution characteristics of the agricultural land confining the River Yamuna in Delhi. Agricultural fields surrounding Yamuna river are the point source of pollution that directly impact the river quality.

Sankhua et al. (2012) focused on concepts of hydraulic flood routing model, with time-varying roughness updating to simulate flows through natural channels, based on the quasi-steady dynamic wave and full dynamic wave theory, emphasizing the solving of the intricate Saint Venant's equations.

Parhi et al. used HEC-RAS to calibrate the channel roughness coefficient (Manning's n value) along the river Mahanadi, Odisha. The authors concluded that Manning's "n" value of 0.032 gives best result for Khairmal to Munduli reach of the Mahanadi river. The calibrated model, in terms of channel roughness, was used to simulate the flood for year 2006 in the same river reach. The performance of the calibrated and validated HEC-RAS based model is tested using Nash and Sutcliffe efficiency

Arunesh (2012) estimated the design flood at Hathnikund and Okhla barrages for different return levels. An analysis of the frequency of flood peaks for different threshold was also carried out at both the barrages. Two methods have been used for the estimation of design flood, namely Gumbel's extreme value distribution and Log Pearson type 3 distribution. The result of frequency analysis clearly indicates that the frequency of extreme events has increased in the recent past. This result is significant as it clearly indicates that there is an urgent need to develop flood prevention and mitigation measures for Delhi and NCR region

Mehta et al., (2012) presented a preliminary design for the physical enhancement of the reach of the Tapi River located near the confluence of Arabian Sea and the Tapi River in

Surat City, Gujarat. The design of table channel was carried out using the Copeland method, which has been included in the HEC-RAS model. A review of the application of remote sensing and GIS in flood management with particular focus on the developing countries of Asia has been presented by Sanyal et al., (2004)

CHAPTER 3

METHODOLOGY

3.1 OVERVIEW

- A digital elevation model (DEM) is a digital file consisting of terrain elevations for ground positions at regularly spaced horizontal intervals.
- Its uses range from scientific, commercial, industrial, operational to military applications.
- In the academic or a research institution, DEM is used primarily as an input or as a data source itself in studies along the fields of climate impact studies, water & wildlife management, geological hydrological modeling, geographic information technology, geomorphology & landscape analysis, mapping purposes, & educational program. However, sometimes DEM is not available

3.2 BASIC CONCEPTS AND DEFINITION

Topographic map - is a type of map characterized by large-scale detail and quantitative representation of relief, usually using contour lines in modern mapping. It is a detailed & accurate graphic representation of cultural & natural features on the ground.

Contour lines—Contour line connects a series of points of equal elevation & is used to illustrate relief on a map. For example, numerous contour lines that are close to one another show hilly or mountainous terrain; when far apart, they indicate a gentler slope.

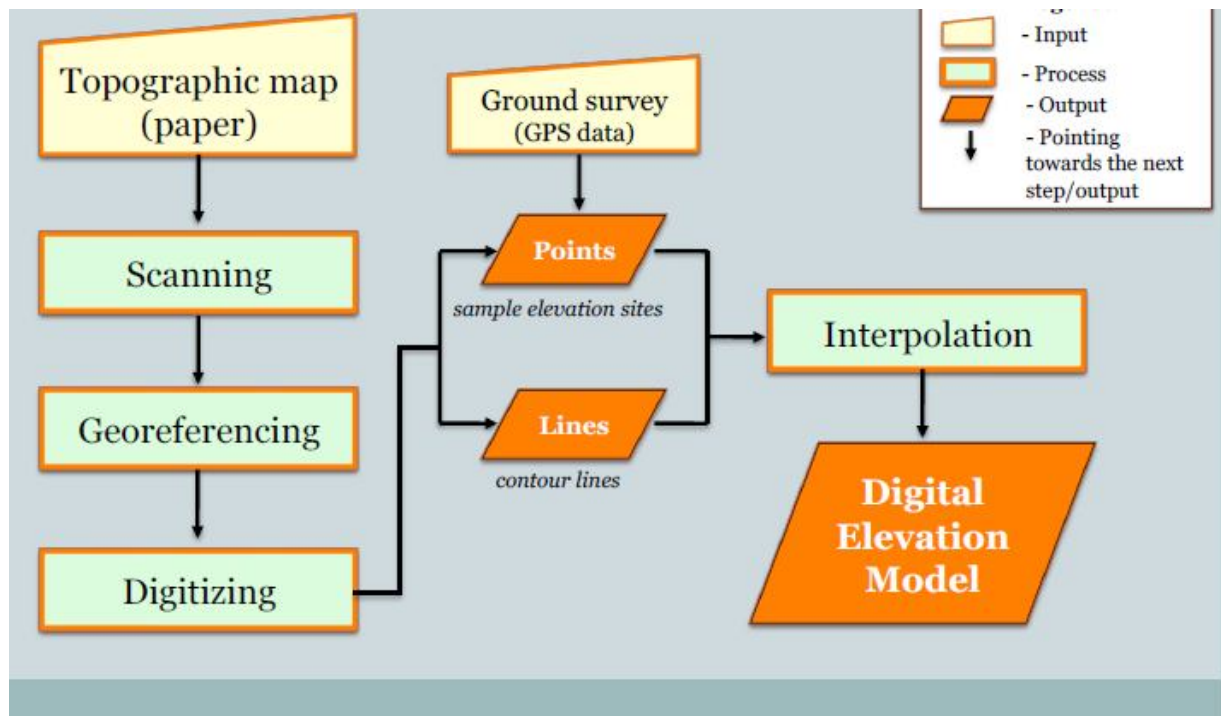
Scanning - is a process of converting any paper-based material (in this case, paper based topographic map) into a digital format, which is usually integrated into the GIS database

Geo referencing - refers to the process of assigning map coordinates to an image data. Data that are already geo referenced can be used as reference in Geo referencing.

Digitizing - is a process of converting spatial features (point, line & polygon) from a paper-based source into a digital form by tracing. This can be done using a digitizing tablet or by on-screen digitizing.

Interpolation - is a process of assigning values to unknown points by using values from a usually scattered set of known points.

3.3 FLOW DIAGRAM FOR CREATING DEM



3.4 UNSTEADY FLOW DATA

In order to perform unsteady flow computations, unsteady flow data are required. Unsteady flow data comprises of-

- Boundary conditions (external as well as internal)
- Initial conditions

3.4.1 BOUNDARY CONDITIONS

Upstream ends of river system can be modeled using these different kind of boundary conditions

- Flow hydrograph
- Stage hydrograph
- Flow and stage hydrograph

Similarly for modeling of downstream end of river system following can be used

- Normal depth (Manning's n equation)
- Rating curve
- Stage hydrograph
- Flow hydrograph
- Flow and stage hydrograph

We can also establish boundary conditions at the internal locations at internal cross sections

- Lateral inflow hydrograph
- Uniform lateral inflow hydrograph
- Groundwater interflow
- Internal stage and flow hydrograph

Gated structure which are defined within in a system like-(inline, lateral, between storage areas and or 2D flow areas) can have various kinds of boundary conditions to control the gates-

- Time series of gate opening
- Elevation controlled gate
- Navigation dams rules
- Internal observed stage and flow

3.4.2 INITIAL CONDITIONS

We have to add initial conditions along with the boundary conditions before doing the simulation at all the nodes in the system. There are two ways to establish initial conditions. Mostly used is adding the flow data for each reach and then have a program to compute steady flow backwater analysis. Another way is to use the previous run by writing a file of flow and stage using previous run, which can be used as initial condition for subsequent run.

STEPS FOLLOWED FOR SIMULATION IN HEC RAS

3.5 IMPORT IN HEC RAS

3.5.1 Start the HEC-RAS project

- Select the location and give name for the project

3.5.2 Establish the Units and Setup

- Options >Units System to confirm the Units(here done using SI units)

3.5.3 HEC-RAS Mapper

- HEC-RAS Opening window>RAS Mapper
- Populate the terrain layer

3.5.4 Assigning the projection

Before creating a terrain, a projection setting is done This is controlled by a projection file, extension .prj.

3.5.5 Add the terrain

- Browse the location for Terrain file having the Projection .tif
- New terrain layer>Create to import the Terrain

3.5.6 Adding an Aerial photo

- With RAS MAPPER we can also add imagery and other Geo-spatial Data
- Map layer>Add map data layers>Browse the .jpg file to add aerial photo
- After importing data into RAS MAPPER it becomes available in the HECRAS Geometry Editor
- Here we can assign flood data along with boundaries, establish model grid size for flood analysis, edit the surface break line, assign different Manning's n value

3.5.7 Creating the 2D flow area

- Click on edit geometric data
- The surface will initially represent colors depicting the elevations
- Select the extent of the terrain as the boundary. This is a well-defined basin, when we select the area we can establish a calculation mesh across the surface.
- Give the name of the total flow area
- In the 2D Flow area we can apply a computational mesh, Establish Manning's n Values to apply to the default area and for any other land cover areas added.
- Generate computation points on regular interval with all break lines
- Set computation point spacing DX and DY both and generate points in 2D flow area.
- The 2D Area break lines command can be used to incorporate break lines into the flow area to better represent valley lines and ridge for the surface.

3.5.8 Setting Boundary conditions (for incoming and outgoing flow)

- Flow comes from the main river and the two tributaries so incoming boundary condition is set across the cross section of these river and the 2 tributaries
- Name the inflow boundary as main river, tributary 1 and tributary 2
- Similarly create the outflow boundary, low side of the surface is set as outflow boundary thus water can exit along this boundary.

3.5.9 Establish flow condition at the boundary

- Copy the spreadsheet inputs tools such as excel to create flow information to HECRAS
- For Out flowing boundary conditions, set for a normal depth and apply a friction slope to describe how exiting flow is handled.
- Edit unsteady flow data> Add boundary condition location> Add flow area to include total flow area
- Storage/2D flow area>Outflow>normal depth to establish Outflow condition along this line
- Type friction slope 0.1>Ok
- Storage/2D flow area>inflow>flood hydrograph to establish a time dependent flow to this inflow line
- For data time interval set the value to 1hr using the pick list
- For EG slope for distributing flow along BC line assign value 0.1
- Save>unsteady flow data>Ok

3.5.10 Re-Establish the Area Condition

Error may be received if one move directly to analyzing the flow, without first returning to the geometric data editor form and Re-establishing the grid calculations to ensure that they are correctly saved following the flow assignment.

- HEC RAS interface>Edit Geometric Data>2D Area >edit 2D flow Area
- Generate computation points on regular interval>accept the default spacing>Generate points in 2D Area>OK>Save Geometry Data.
- Close the Geometry Editor

3.6 ANALYSIS IN HECRAS

- Set up time for calculating the flow, Reporting time and total analysis duration.

- RAS Mapper is then used to review flood extents over the time, assess the peak extents, flow depths and flow velocities.

3.6.1 Check boxes -

- I. Geometry Preprocessor
- II. Unsteady flow simulation
- III. Post Processor

3.6.2 Set the simulation Time Window

- I. Pick a start date and end date
- II. Set the starting date to 0:00
- III. Set the ending time to 24:00

3.6.3 Set the computational Settings

- I. Computational Interval - 1 min.
- II. Hydrograph Output interval - 1 hour
- III. Mapping Output interval - 1 hour
- IV. Detailed Output interval - 1 hour

We can view the results directly inside HECRAS via RAS MAPPER, we can also generate the flood of the extents and other flow outcomes and export this to a Geo spatial file format

3.7 FLOW DATA TABLES

3.7.1 Flow data for Main River

Date/time	Discharge(m ³ /s)	Date/time	Discharge(m ³ /s)	Date/time	Discharge(m ³ /s)
10/12/2010; ;1800	9.8	;0300	107.7	;1200	871.4
;1900	9.7	;0400	141.7	;1300	839.7
;2000	9.6	;0500	181.6	;1400	808.7
;2100	9.4	;0600	224.8	;1500	778.1
;2200	9.2	;0700	266.4	;1600	748
;2300	8.9	;0800	304	;1700	718.7
;2400	8.6	;0900	337.8	;1800	690.8
11/12/2010 ;0100	8.3	;1000	370.1	;1900	664.1
;0200	8.1	;1100	404.1	;2000	638.5
;0300	7.8	;1200	442.8	;2100	614.3
;0400	6.8	;1300	490.1	;2200	591.3
;0500	7.2	;1400	550.6	;2300	569.1
;0600	7	;1500	625.8	;2400	547.9
;0700		;1600		14/12/2010 ;0100	527.6
	6.8		725.4	;0200	508
;0800	6.8	;1700	836.8	;0300	489.1
;0900	7	;1800	952.2	;0400	471
;1000	7.6	;1900	1061.4	;0500	453.6
;1100	8.6	;2000	1159.4	;0600	436.9
;1200	10.2	;2100	1245.3	;0700	420.7
;1300	12.5	;2200	1319.3	;0800	405.1
;1400	15.5	;2300	1376.9	;0900	390.1
;1500	19.1	;2400	1407.7		
;1600		13/12/2010 ;0100		;1000	
	23.2		1407.8		375.6
;1700	27.5	;0200	1376	;1100	361.7
;1800	31.7	;0300	1320.2	;1200	348.3
;1900	35.6	;0400	1256.4	;1300	348.3
;2000	38.9	;0500	1191.1	;1400	335.4
;2100	41.6	;0600	1126.9	;1500	311
;2200	44.4	;0700	1067.7	;1600	299.5
;2300	48.4	;0800	1017.1	;1700	285.8
;2400	54.9	;0900	974.8	;1800	241.6
12/12/2010 ;0100	65.7	;1000	937.8	;1900	240
;0200	82.7	;1100	904	;2000	239

3.7.2 Flow data for Tributary 1

Date/time	Discharge(m ³ /s)	Date/time	Discharge(m ³ /s)	Date/time	Discharge
10/12/2010; ;1800	0.98	;0300	10.77	;1200	83.97
;1900	0.97	;0400	14.14	;1300	80.87
;2000	0.94	;0500	18.16	;1400	77.81
;2100	0.92	;0600	22.48	;1500	74.8
;2200	0.89	;0700	26.64	;1600	71.87
;2300	0.86	;0800	30.4	;1700	69.08
;2400	0.83	;0900	33.78	;1800	66.41
11/12/2010 ;0100	0.81	;1000	37.06	;1900	63.85
;0200	0.78	;1100	40.41	;2000	61.43
;0300	0.75	;1200	44.28	;2100	59.13
;0400	0.72	;1300	49.01	;2200	56.91
;0500	0.7	;1400	55.06	;2300	54.79
;0600	0.7	;1500	62.88	;2400	52.76
;0700	0.68	;1600	72.54	14/12/2010 ;0100	50.8
;0800	0.68	;1700	83.68	;0200	48.91
;0900	0.7	;1800	95.22	;0300	47.1
;1000	0.76	;1900	106.14	;0400	45.36
;1100	0.86	;2000	115.94	;0500	43.69
;1200	1.02	;2100	124.33	;0600	42.07
;1300	1.25	;2200	131.93	;0700	40.51
;1400	1.55	;2300	137.66	;0800	39.01
;1500	1.91	;2400	140.77	;0900	37.56
;1600	2.32	13/12/2010 ;0100	140.77	;1000	36.17
;1700	2.75	;0200	137.6	;1100	36.17
;1800	3.17	;0300	132.02	;1200	34.83
;1900	3.56	;0400	125.64	;1300	33.54
;2000	3.89	;0500	119.11	;1400	31.1
;2100	4.16	;0600	112.69	;1500	29.95
;2200	4.44	;0700	106.77	;1600	28.58
;2300	4.84	;0800	101.71	;1700	24.16
;2400	5.48	;0900	97.48	;1800	24
12/12/2010 ;0100	6.57	;1000	93.78	;1900	23.9
;0200	8.27	;1100	90.4	;2000	

TABLE 2

3.7.2 Flow Data for tributary 2

Date/time	Discharge(m ³ /s)	Date/time	Discharge(m ³ /s)	Date/time	Discharge(m ³ /s)
10/12/2010		;0300		;1200	
;1800	0.49		7.07		43.57
;1900	0.485	;0400	9.08	;1300	41.985
;2000	0.48	;0500	11.24	;1400	40.435
;2100	0.47	;0600	13.32	;1500	38.905
;2200	0.46	;0700	15.2	;1600	37.4
;2300	0.445	;0800	16.89	;1700	35.935
;2400	0.43	;0900	18.505	;1800	34.54
11/12/2010		;1000		;1900	
;0100	0.415		20.205		33.205
;0200	0.405	;1100	22.14	;2000	31.925
;0300	0.39	;1200	22.14	;2100	30.715
;0400	0.375	;1300	24.505	;2200	29.565
;0500	0.36	;1400	27.53	;2300	28.455
;0600	0.35	;1500	31.44	;2400	27.395
;0700		;1600		14/12/2010	
	0.375		36.27	;0100	26.38
;0800	0.36	;1700	41.84	;0200	25.4
;0900	0.35	;1800	47.61	;0300	24.455
;1000	0.38	;1900	53.07	;0400	23.55
;1100	0.43	;2000	57.97	;0500	22.68
;1200	0.51	;2100	62.265	;0600	21.845
;1300	0.625	;2200	65.965	;0700	21.035
;1400	0.775	;2300	68.83	;0800	20.225
;1500	0.955	;2400	70.385	;0900	19.505
;1600		13/12/2010		;1000	
	1.16	;0100	70.39		18.78
;1700	1.375	;0200	68.8	;1100	18.085
;1800	1.585	;0300	66.01	;1200	17.415
;1900	1.78	;0400	62.82	;1300	16.77
;2000	1.945	;0500	59.555	;1400	16.15
;2100	2.08	;0600	56.345	;1500	15.55
;2200	2.22	;0700	53.385	;1600	14.975
;2300	2.745	;0800	50.855	;1700	14.29
;2400	3.285	;0900	48.74	;1800	12.08
12/12/2010		;1000		;1900	
;0100	3.285		46.89		12
;0200	5.385	;1100	45.2	;2000	11.95

TABLE 3

3.7.4 RIVER LENGTH AND PROFILES

	Main river	Tributary 1	Tributary 2
Reach length	54.4 km	17.1km(up to confluence with main river)	12.7km(upto confluence with main river)
Channel's Manning's n value	.045-.07	.04-.06	.025-.05
Average discharge(m ³ /s) (100year)	906	130	70

TABLE 4

3.8 RIVER PROFILE

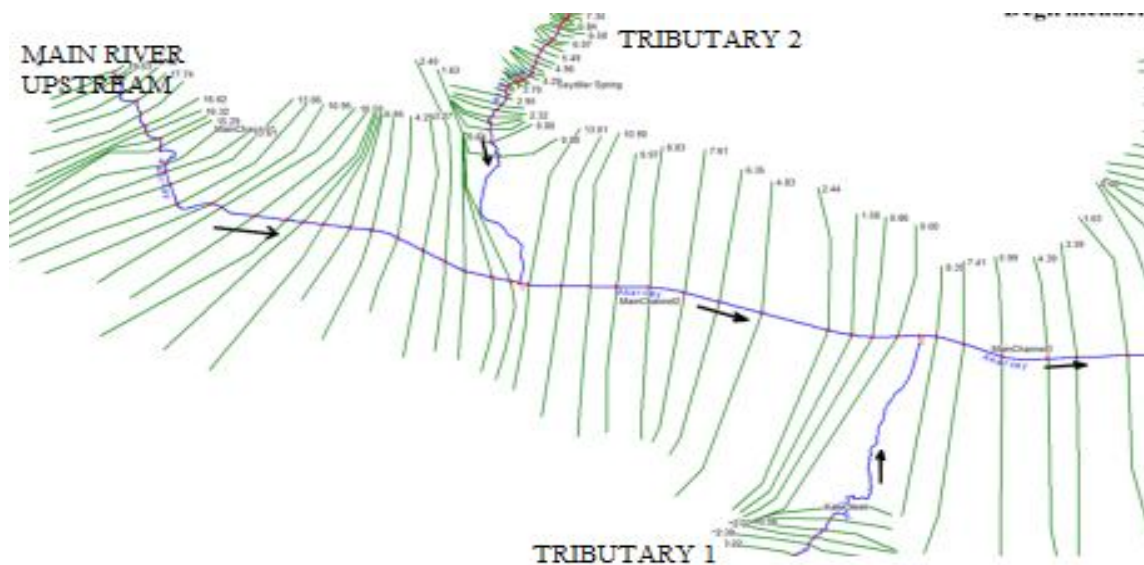


FIG 3.8 – SHOWING RIVER PROFILE IN HEC RAS

3.9

Known Normal depth and Water surface elevation (m)	MAIN RIVER	TRIBUTARY 1	TRIBUTARY 2
100-year	72.05	82.84	72.02
100-year	74.28	83.27	73.74.

TABLE 5

CHAPTER 4

RESULTS

4.1 HEC RAS SOFTWARE

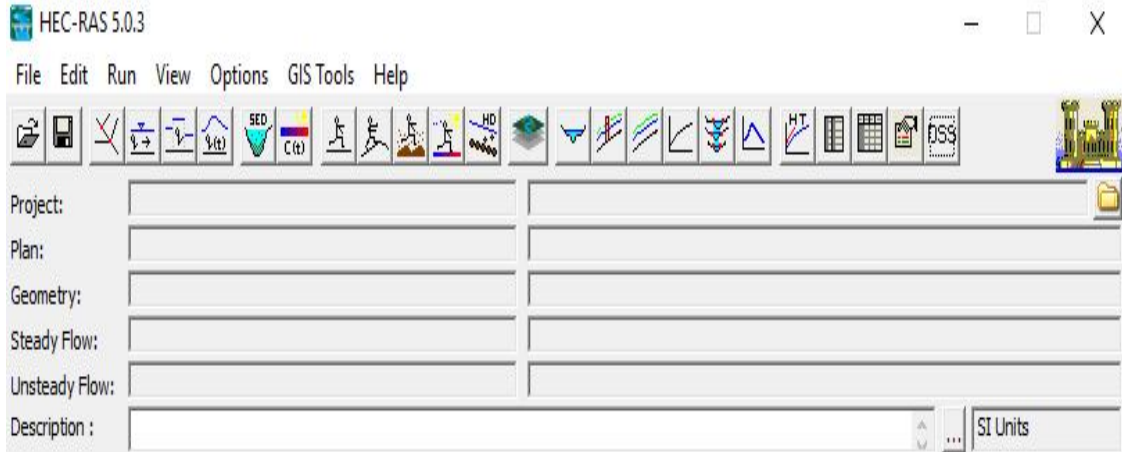


Fig 4.1 HEC RAS OPENING WINDOW

4.2 RAS MAPPER

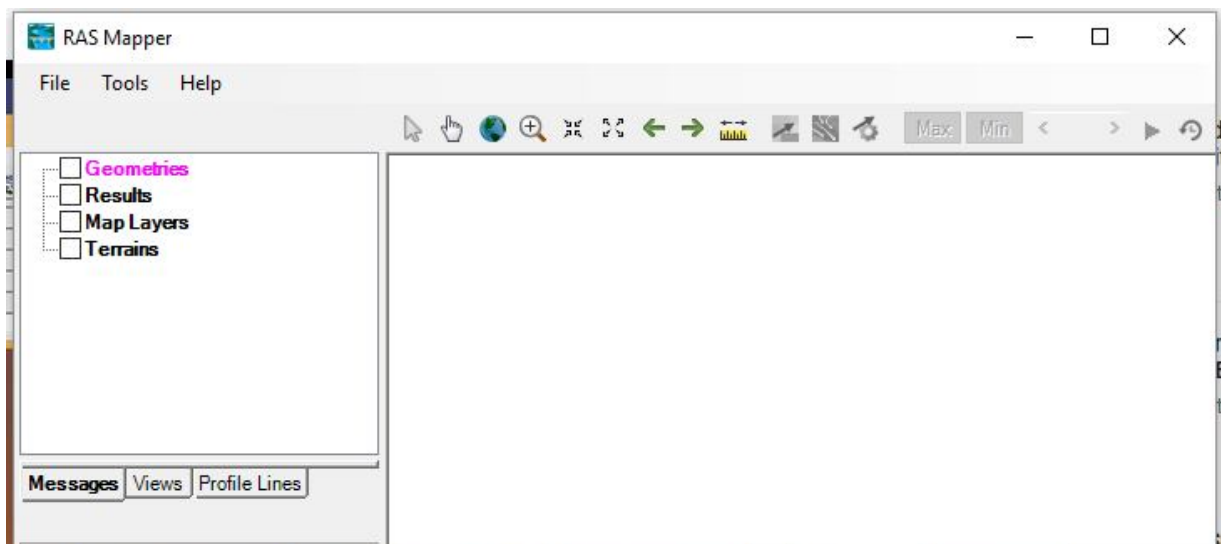


Fig 4.2 RAS MAPPER OPENING WINDOW

4.3 SETTING THE PROJECTION FOR PROJECT

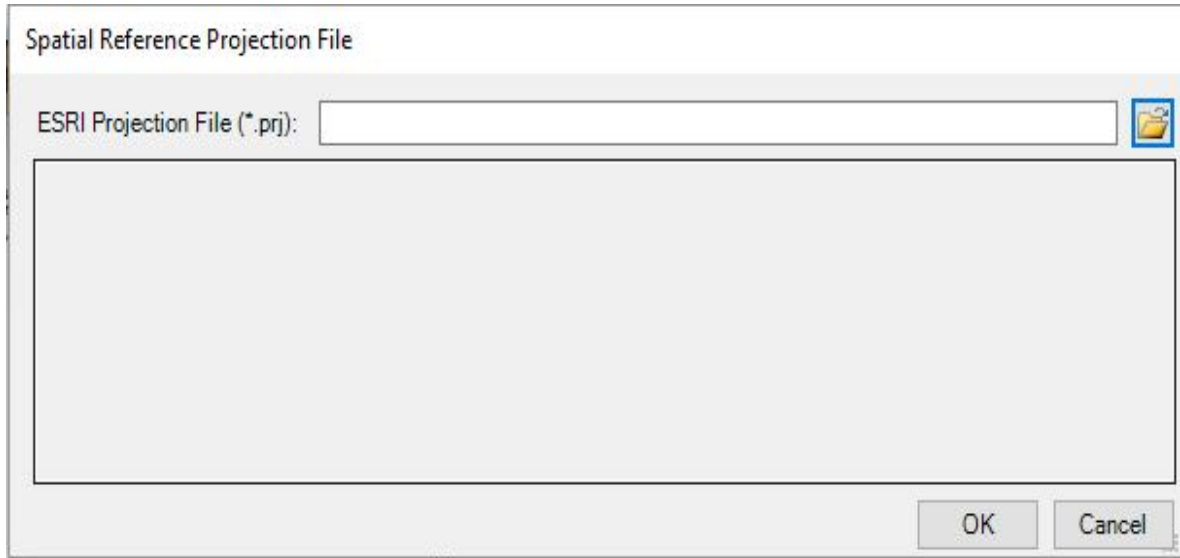


Fig 4.3 SPATIAL REFERENCE PROJECTION FILE WINDOW

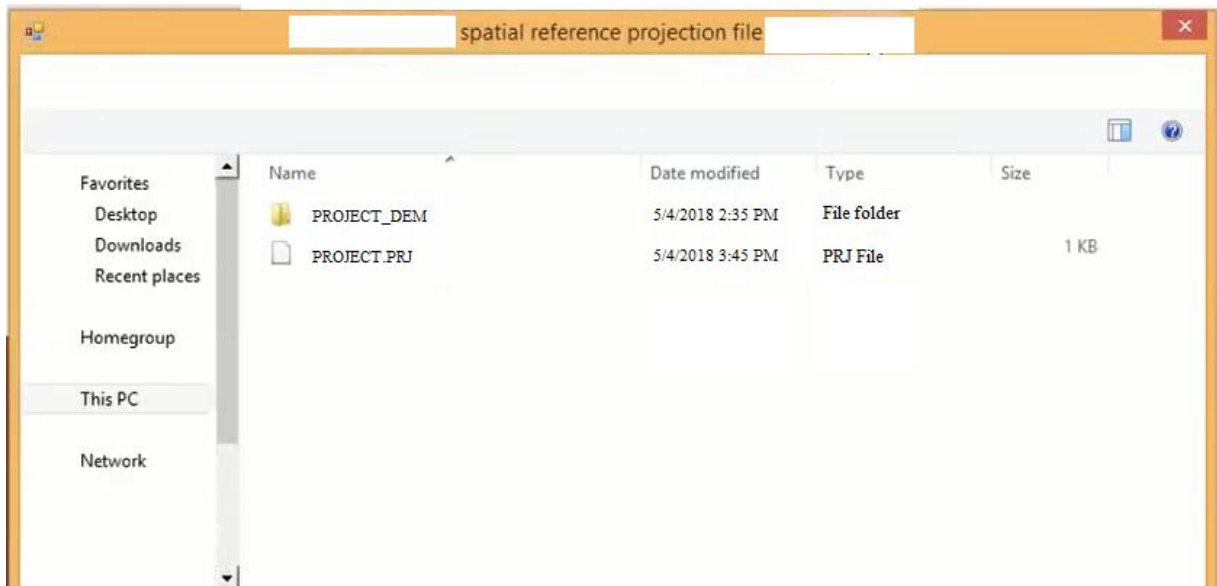


Fig 4.4 SETTING THE PROJECTION FILE

4.4 EXPORTING TERRAIN TO HEC RAS

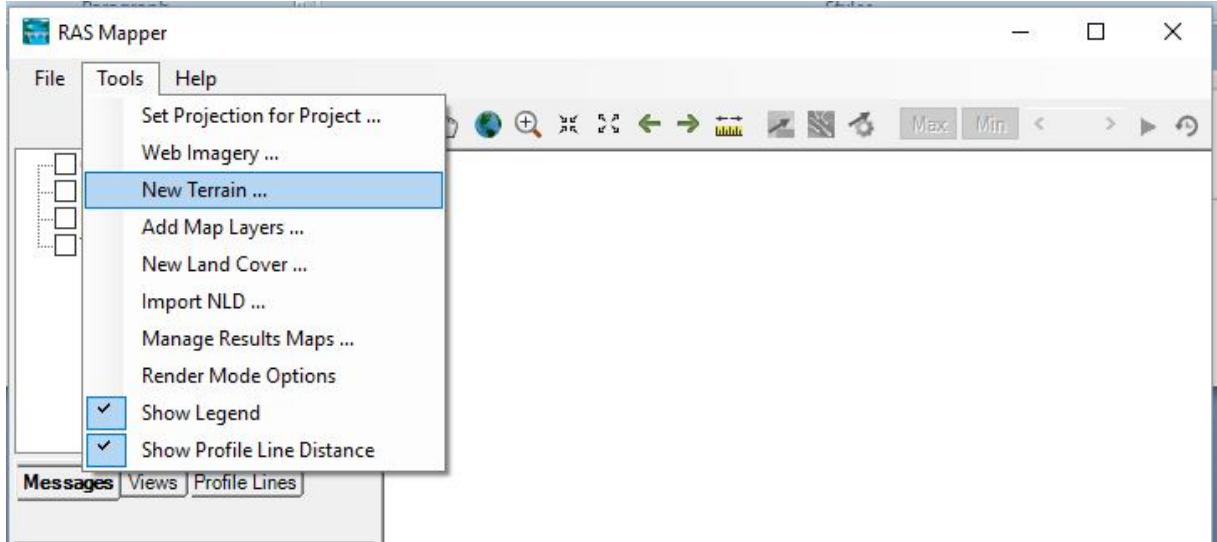


Fig 4.5 LOADING TERRAIN DATA IN RAS MAPPER

4.5 SETTING PRECISION

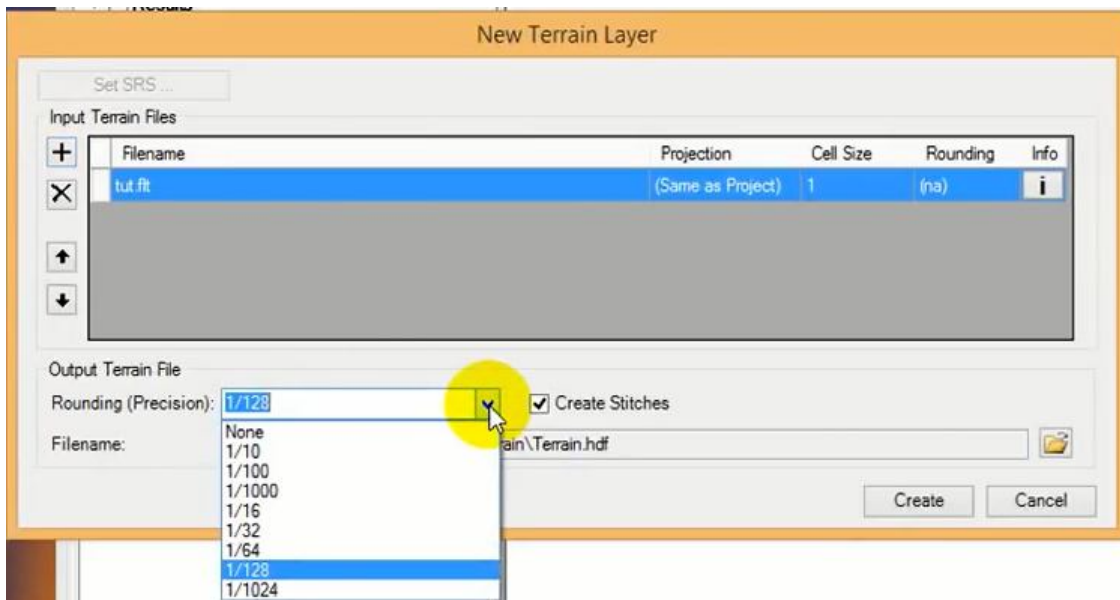


Fig 4.6 SETTING TERRAIN WITH PRECISION

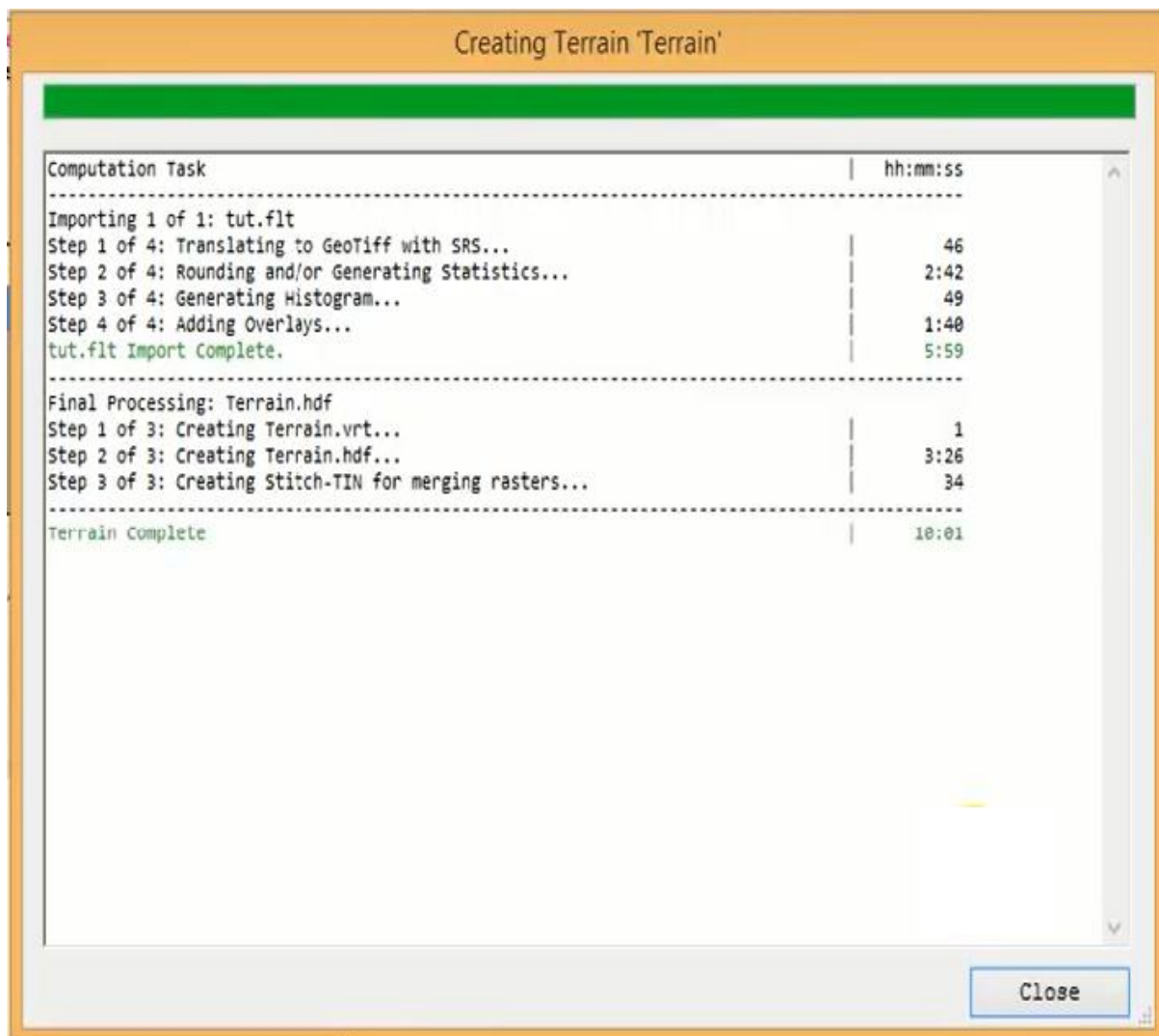


Fig 4.7 TERRAIN LOADING COMPLETED

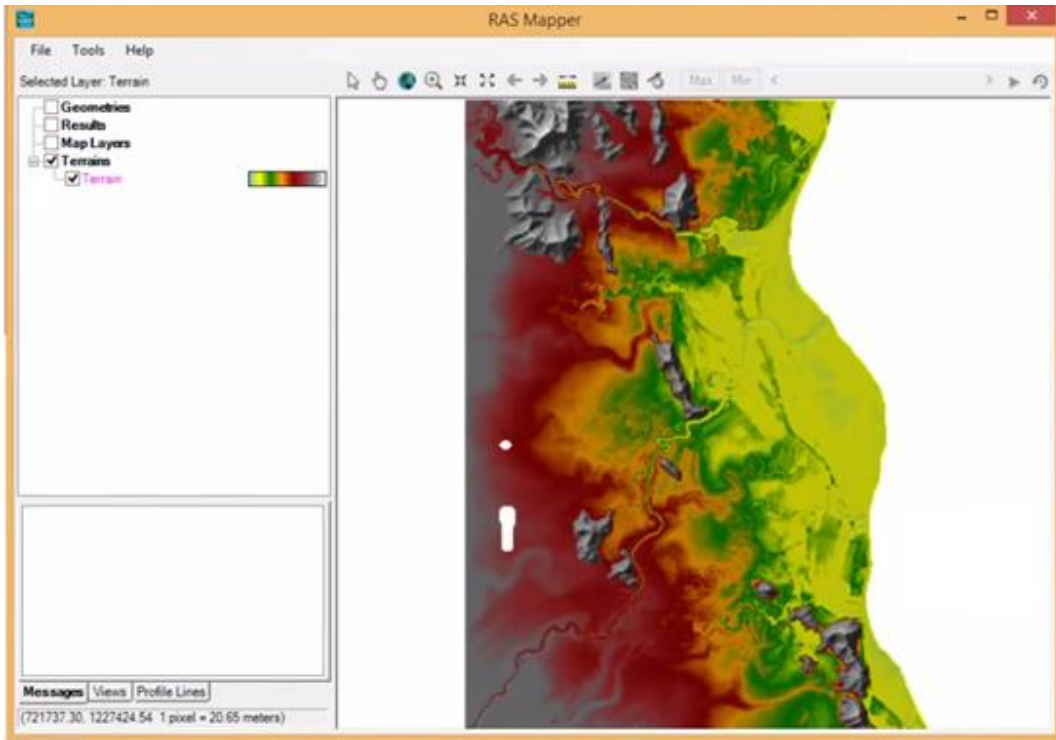


Fig 4.8 TERRAIN IN RAS MAPPER

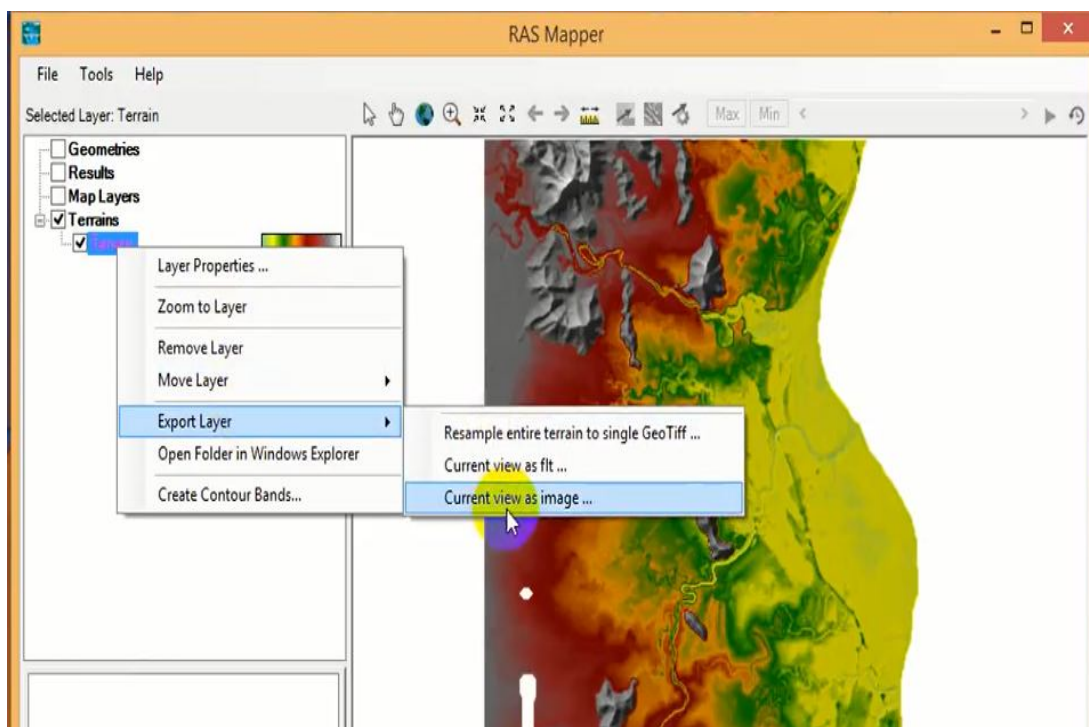


Fig 4.9 CREATE IMAGE OF THE LOADED TERRAIN

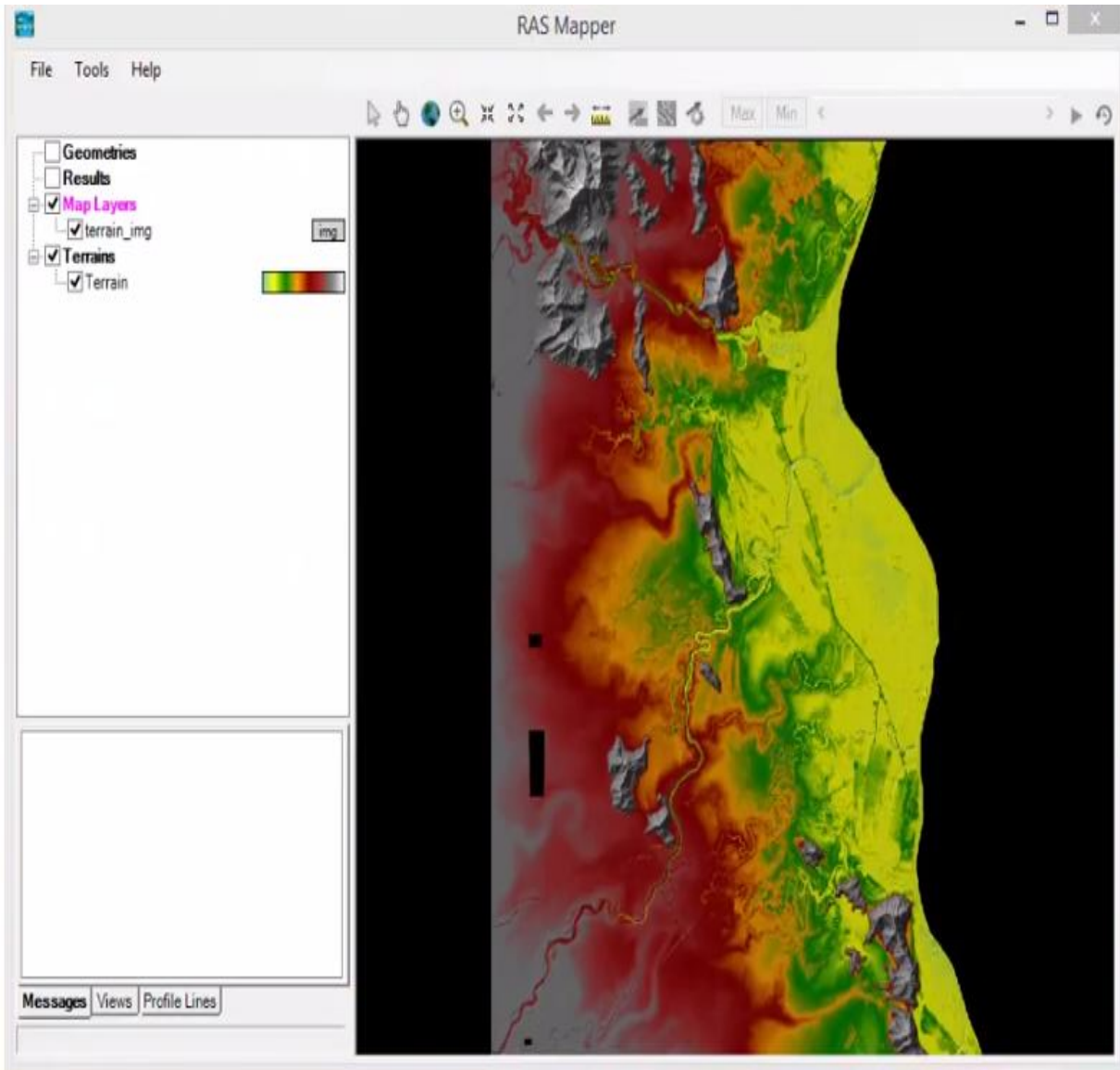


Fig 4.10 IMAGE OF TERRAIN LOADED IN RAS MAPPER

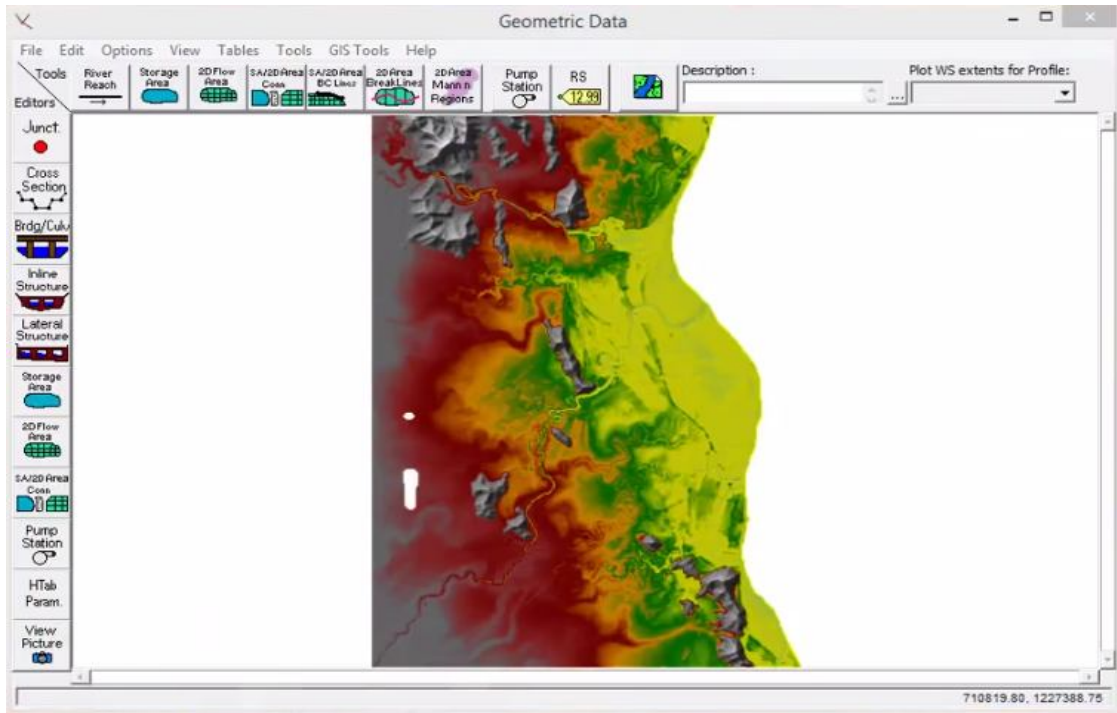


Fig 4.11 GEOMETRIC DATA WINDOW

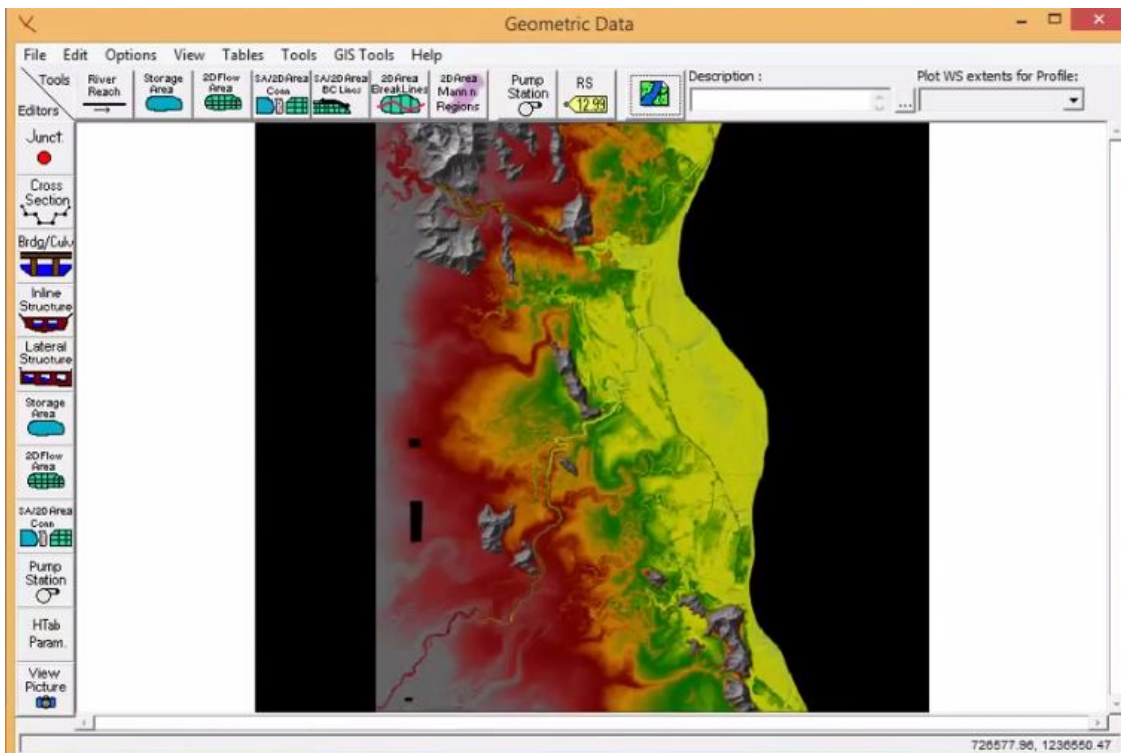


Fig 4.12 CHECKING FOR IMAGE IN GEOMETRIC DATA WINDOW

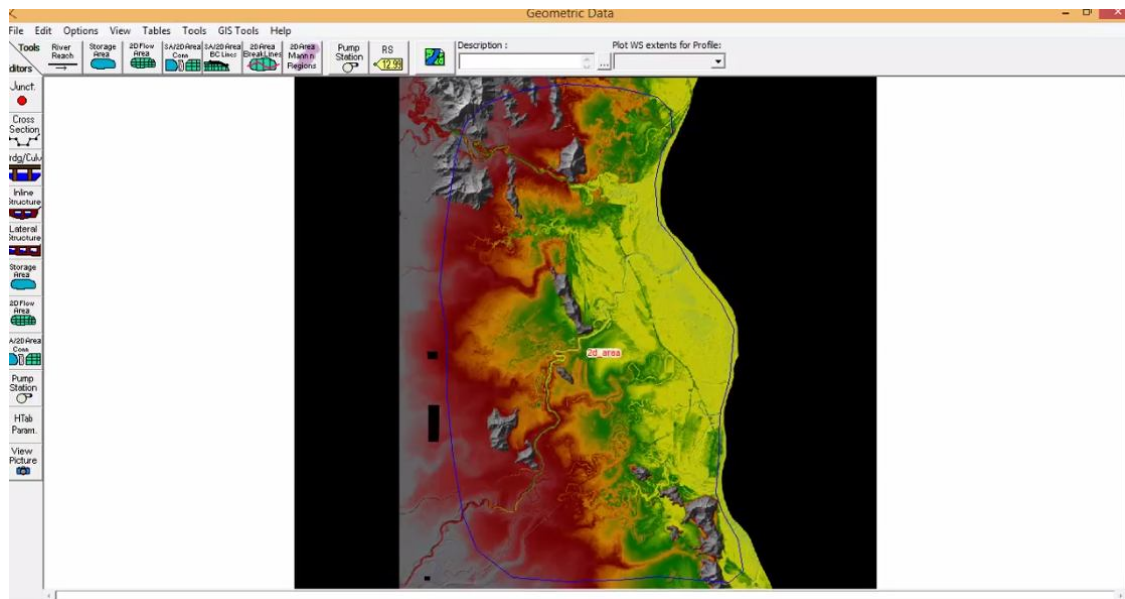


Fig 4.13 2D AREA/ POLYGON CREATED OF STUDY AREA

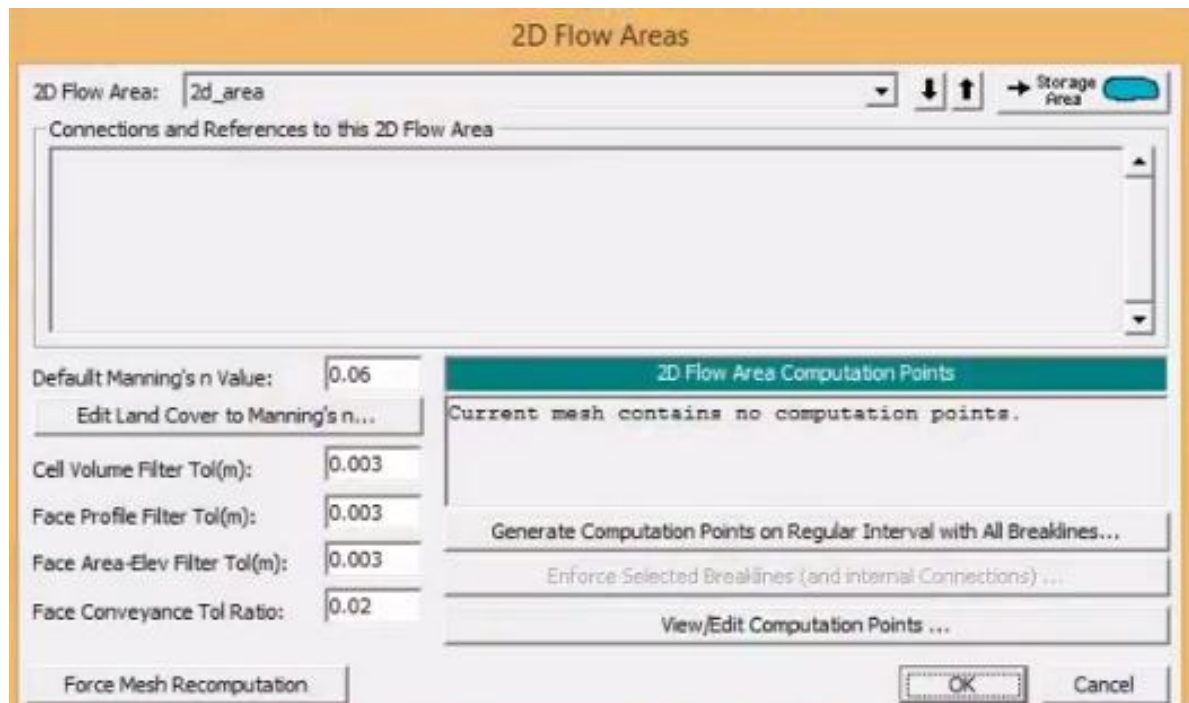


Fig 4.14 MESHING OF 2D FLOW AREA

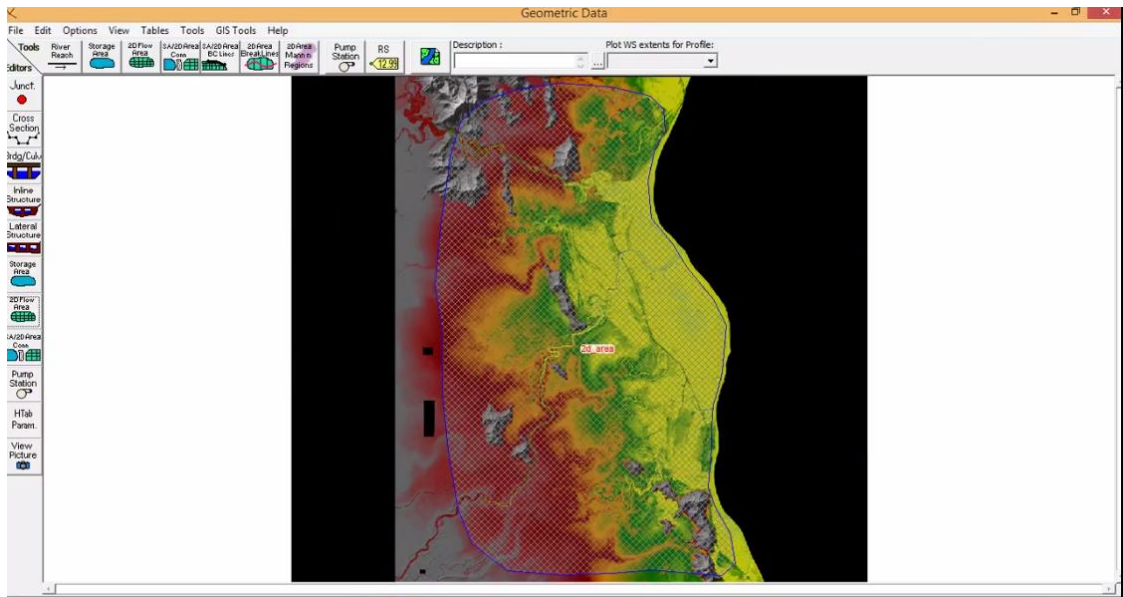


Fig 4.15 MESHING OF STUDY AREA

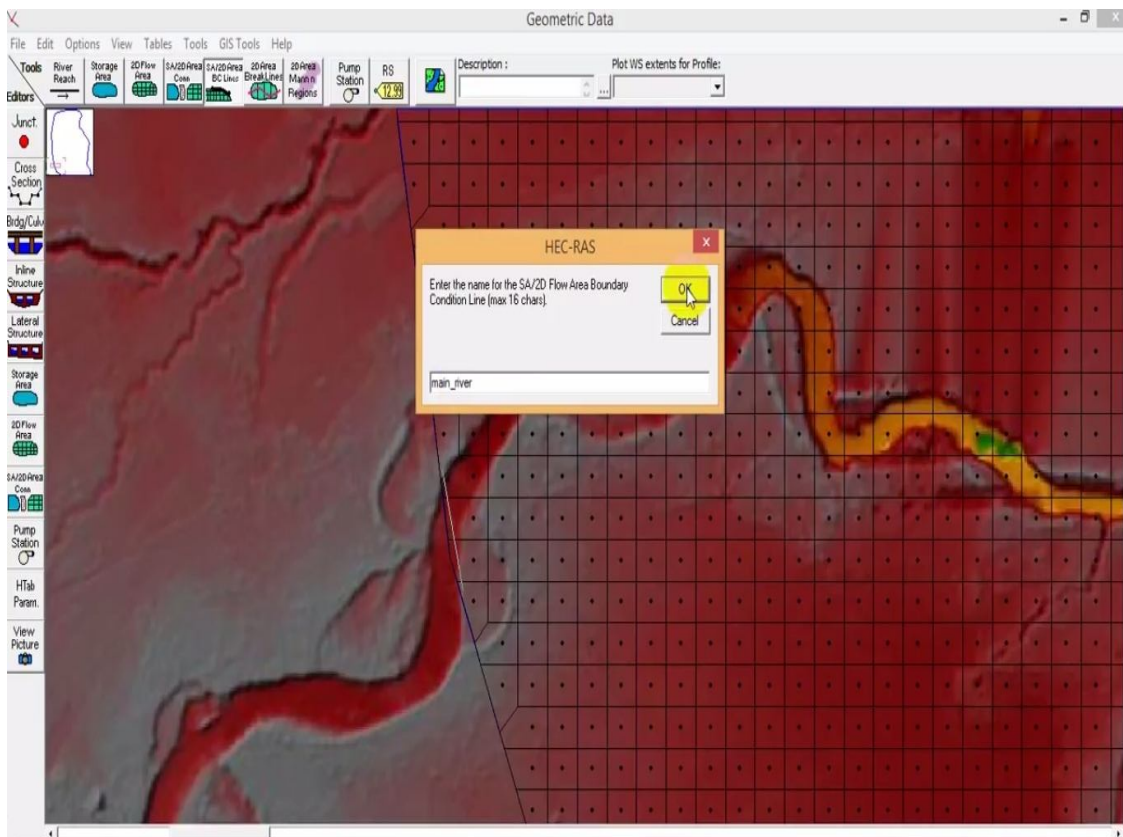


Fig 4.16 DRAWING INLET BOUNDARY CONDITION SHOWING MAIN RIVER

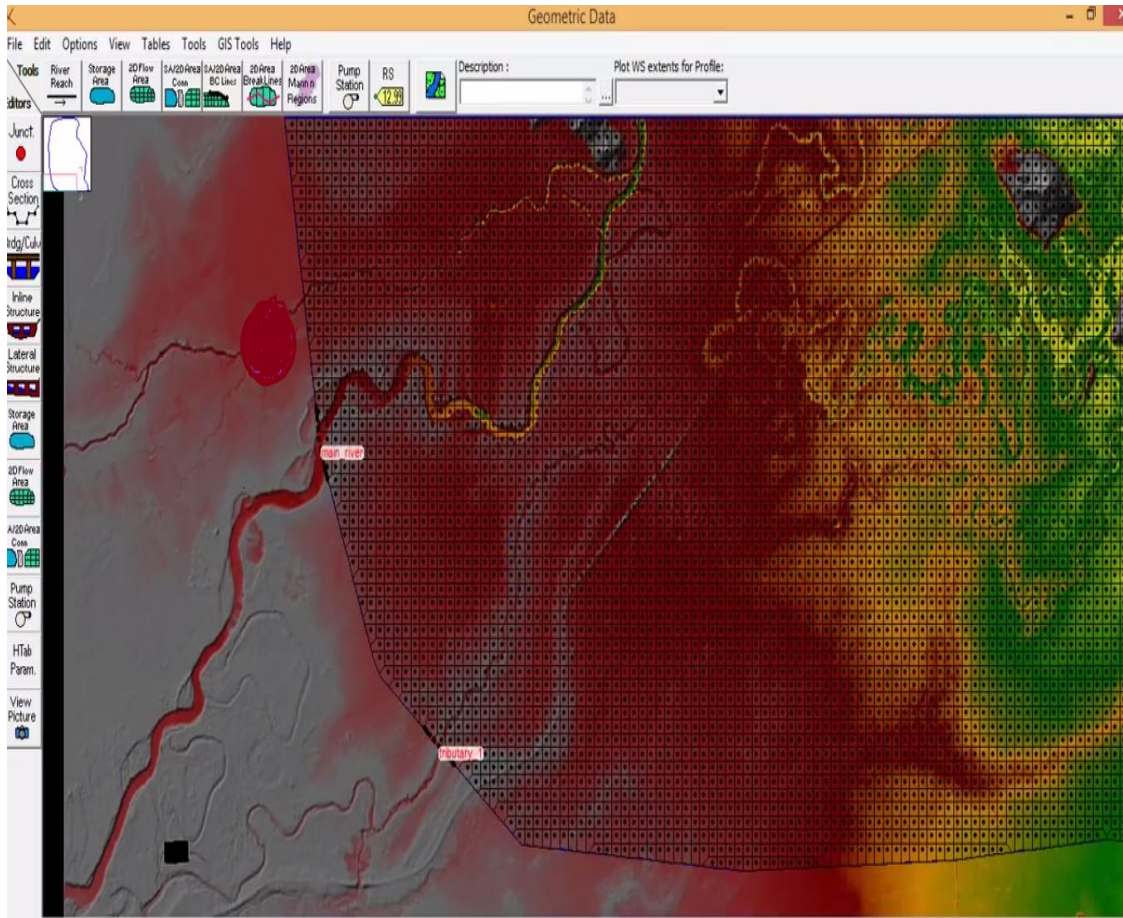


Fig 4.17 MAIN RIVER AND TRIBUTARY 1

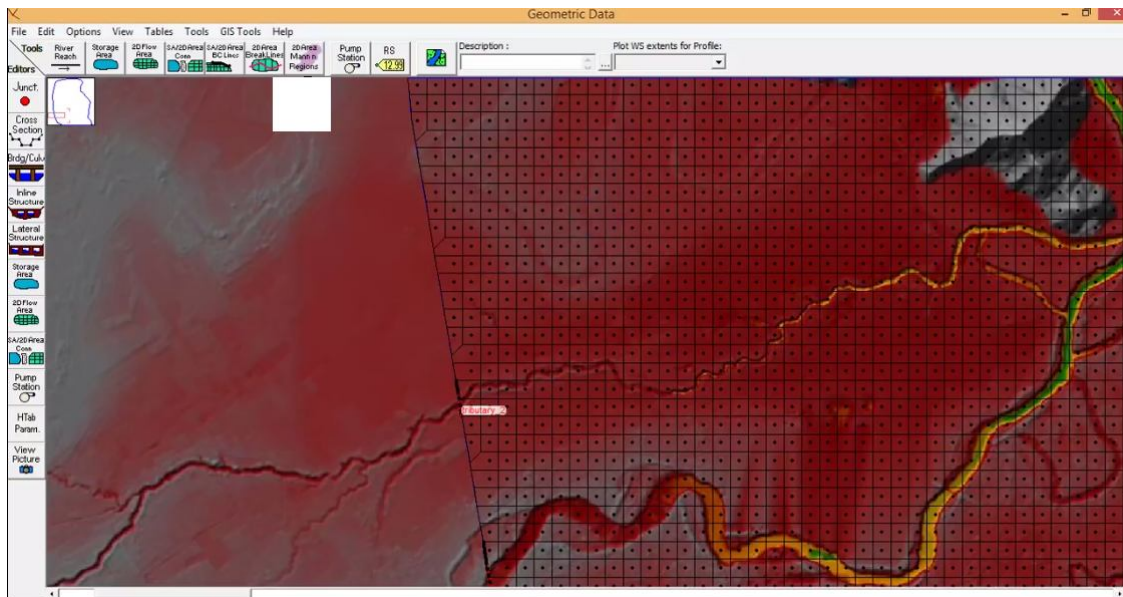


Fig 4.18 TRIBUTARY 2

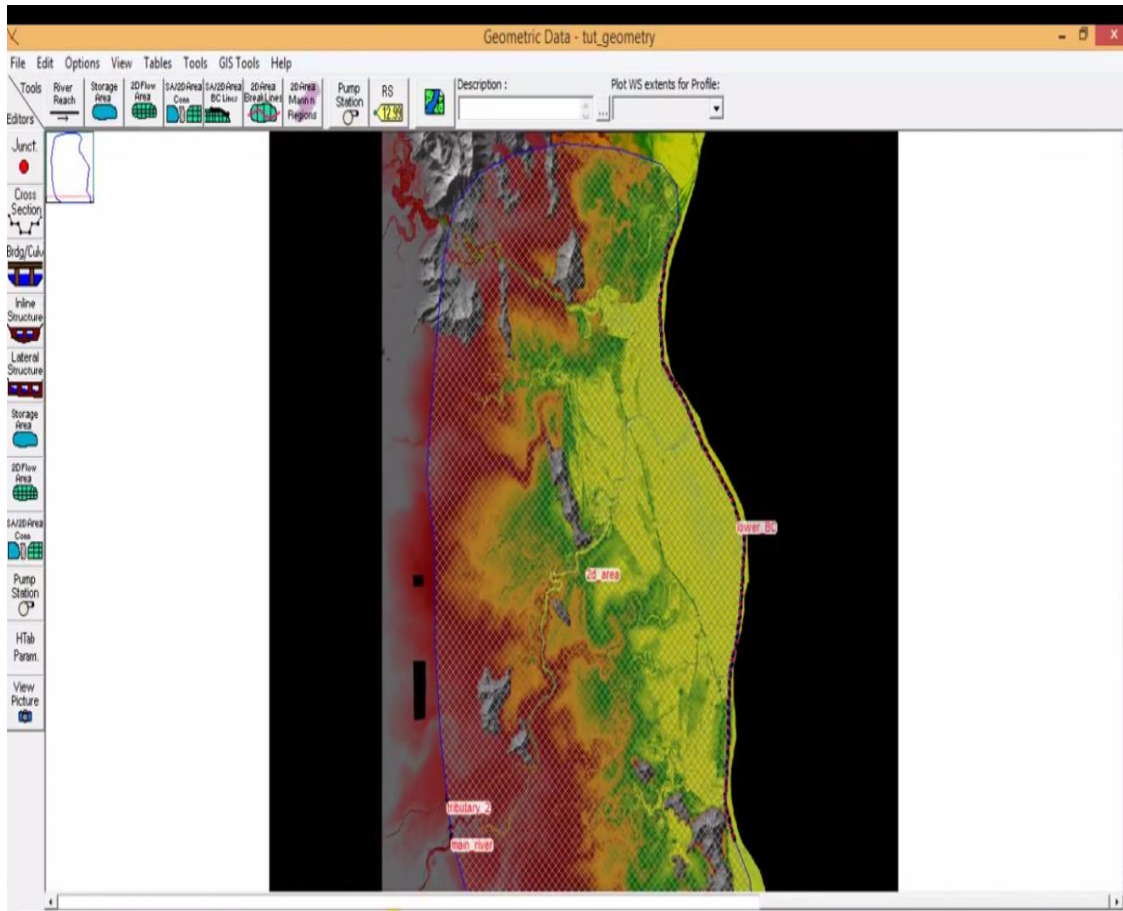


Fig 4.19 INLET, 2D AREA LOWER BOUNDARY CONDITION

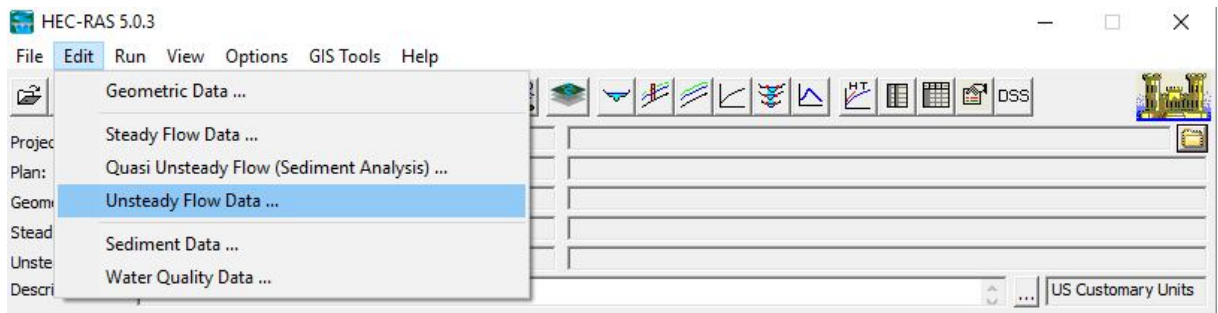


Fig 4.20 ADDING UNSTEADY FLOW DATA

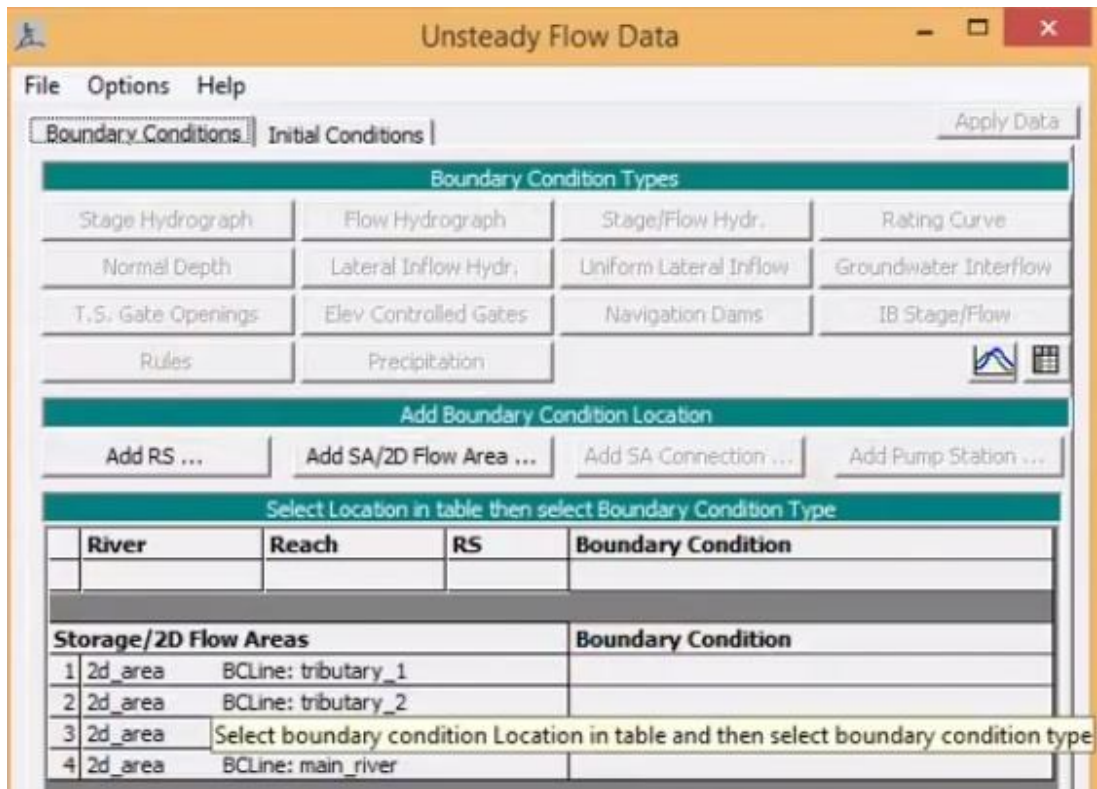


Fig 4.21 ADDING BOUNDARY CONDITION LOCATION AND TYPE

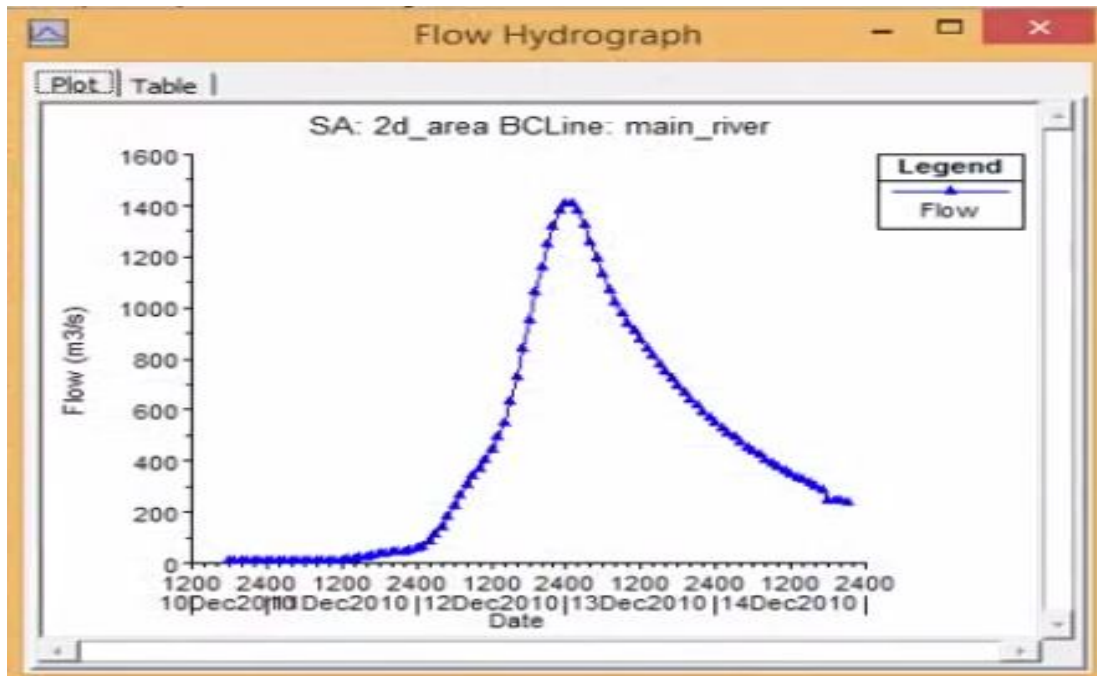


Fig 4.22 FLOW HYDROGRAPH OF MAIN RIVER

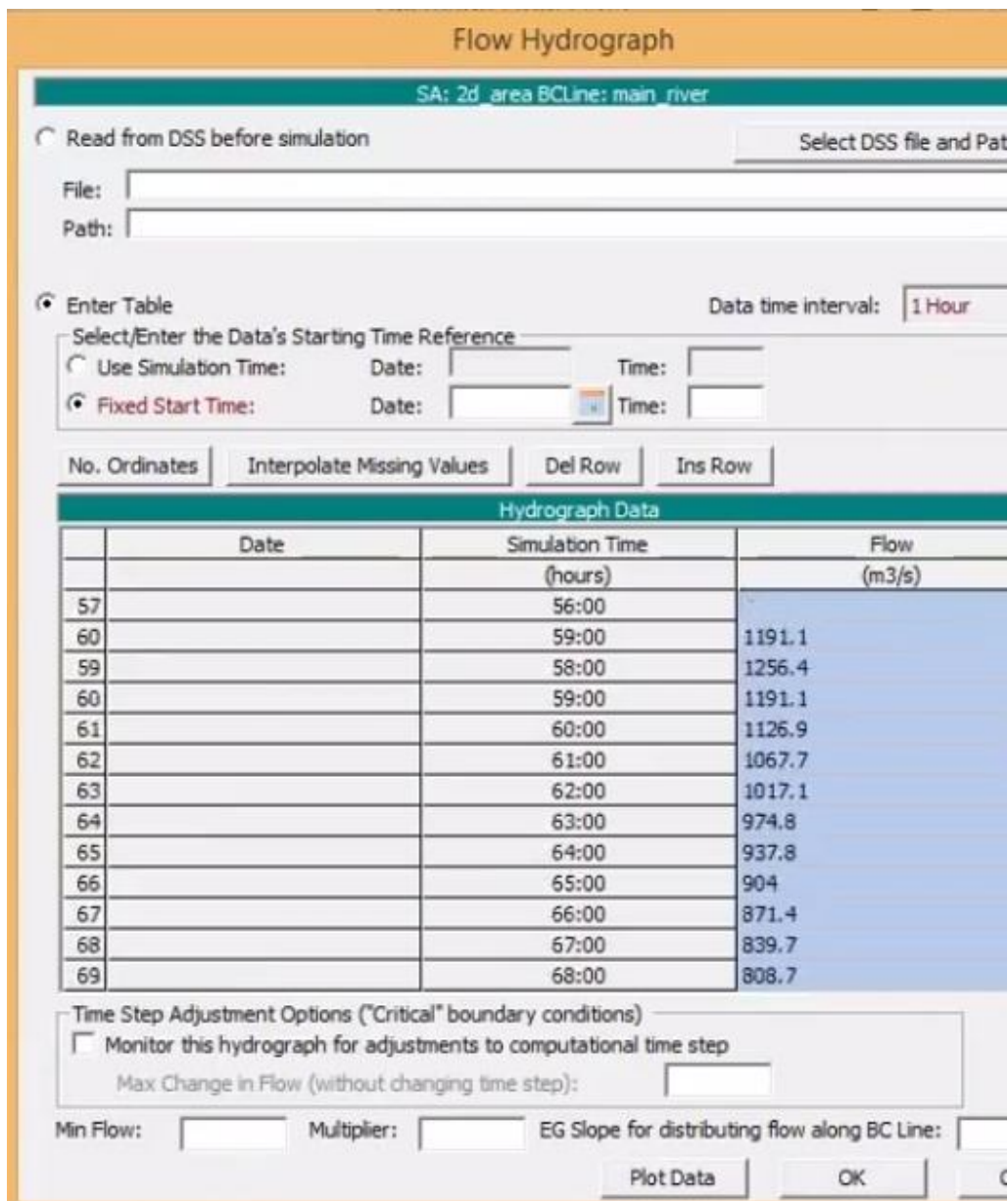


Fig 4.23 FLOW HYDROGRAPH DATA OF MAIN RIVER

Flow Hydrograph

SA: 2d_area BCLine: tributary_1

Read from DSS before simulation Select DSS file and Path

File:

Path:

Enter Table Data time interval: 1 Hour

Select/Enter the Data's Starting Time Reference

Use Simulation Time: Date: Time:

Fixed Start Time: Date: 10DEC2010 Time: 1800

No. Ordinates | Interpolate Missing Values | Del Row | Ins Row

Hydrograph Data			
	Date	Simulation Time (hours)	Flow (m3/s)
88	14Dec2010 0900	87:00	39.01
89	14Dec2010 1000	88:00	37.56
90	14Dec2010 1100	89:00	36.17
91	14Dec2010 1200	90:00	34.83
92	14Dec2010 1300	91:00	33.54
93	14Dec2010 1400	92:00	32.3
94	14Dec2010 1500	93:00	31.1
95	14Dec2010 1600	94:00	29.95
96	14Dec2010 1700	95:00	28.58
97	14Dec2010 1800	96:00	24.16
98	14Dec2010 1900	97:00	24
99	14Dec2010 2000	98:00	23.9
100	14Dec2010 2100	99:00	23.8

Time Step Adjustment Options ("Critical" boundary conditions)

Monitor this hydrograph for adjustments to computational time step

Max Change in Flow (without changing time step):

Min Flow: Multiplier: EG Slope for distributing flow along BC Line:

Plot Data | OK | Cancel

Fig 4.24 FLOW HYDROGRAPH DATA OF TRIBUTARY 10

Flow Hydrograph

SA: 2d_area BCLine: tributary_2

Read from DSS before simulation Select DSS file and Path

File:

Path:

Enter Table Data time interval: 1 Hour

Select/Enter the Data's Starting Time Reference

Use Simulation Time: Date: Time:

Fixed Start Time: Date: 10DEC2010 Time: 1800

No. Ordinates

Hydrograph Data			
	Date	Simulation Time (hours)	Flow (m3/s)
1	10Dec2010 1800	00:00	0.49
2	10Dec2010 1900	01:00	0.485
3	10Dec2010 2000	02:00	0.48
4	10Dec2010 2100	03:00	0.47
5	10Dec2010 2200	04:00	0.46
6	10Dec2010 2300	05:00	0.445
7	10Dec2010 2400	06:00	0.43
8	11Dec2010 0100	07:00	0.415
9	11Dec2010 0200	08:00	0.405
10	11Dec2010 0300	09:00	0.39
11	11Dec2010 0400	10:00	0.375
12	11Dec2010 0500	11:00	0.36
13	11Dec2010 0600	12:00	0.35

Time Step Adjustment Options ("Critical" boundary conditions)

Monitor this hydrograph for adjustments to computational time step

Max Change in Flow (without changing time step):

Min Flow: Multiplier: EG Slope for distributing flow along BC Line:

Fig 4.25 FLOW HYDROGRAPH DATA OF TRIBUTARY 2

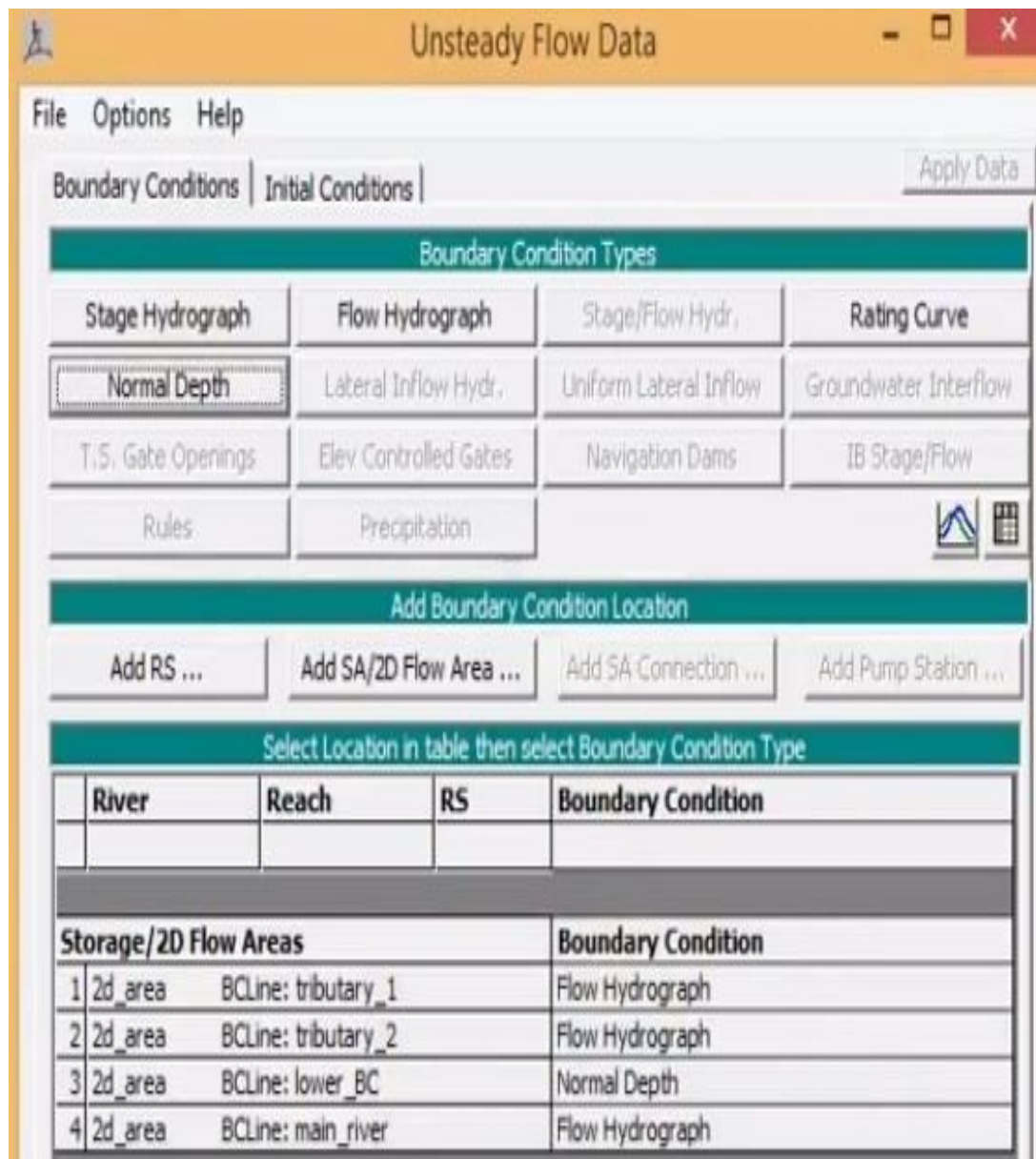


Fig 4.26 ADDING NORMAL DEPTH AS LOWER BOUNDARY CONDITION

Unsteady flow data added for tributary 1 is flow hydrograph at a section

Unsteady flow data added for tributary 2 is flow hydrograph at a section

Unsteady flow data added for main river is flow hydrograph flow at a section

Lower boundary condition is taken as Normal depth

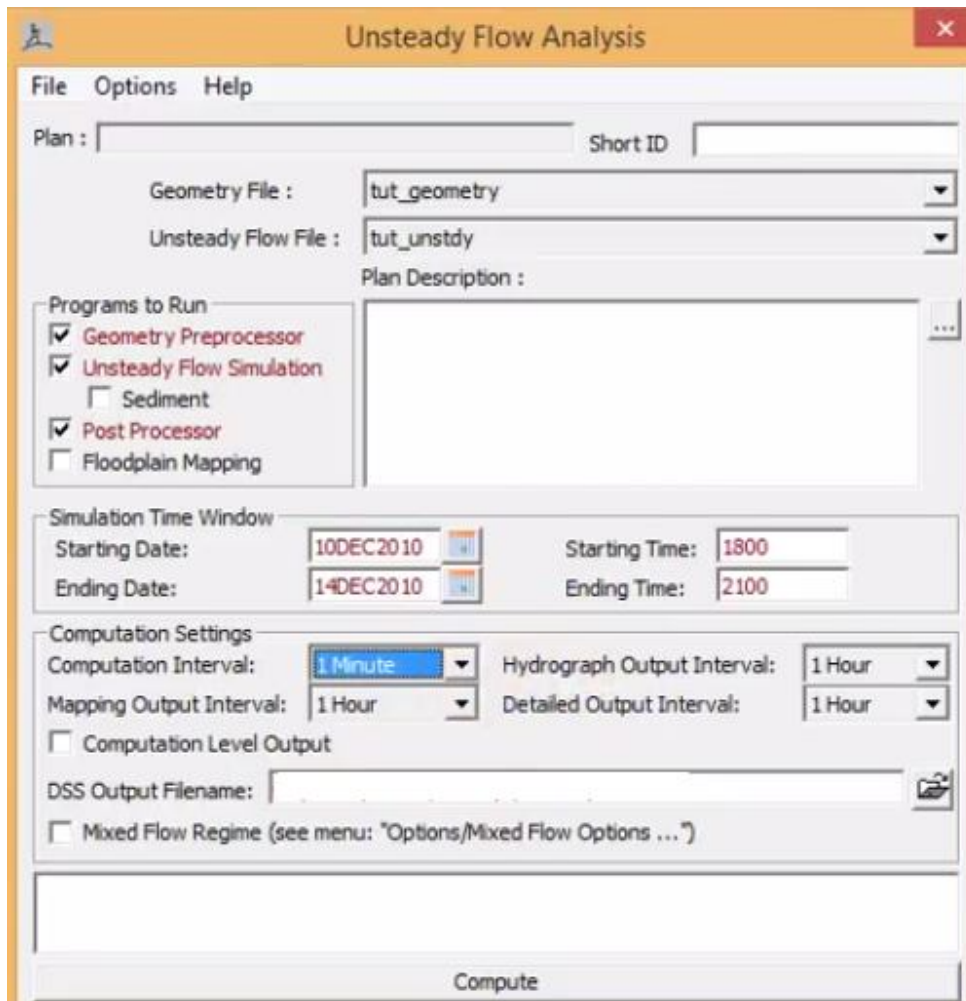


Fig 4.27 UNSTEADY FLOW ANALYSIS

- Simulation time is from taken from 1800 hOur (10 Dec 2010) to 2100 hOur (14 Dec 2010)
- Computation interval is taken as 10 min.
- Mapping Output interval is taken as 1 hOur
- Hydrograph Output interval is taken as 1 hOur
- Detailed Output interval is taken as 1 hOur
- Unsteady flow data added for tributary 1 is flow hydrograph

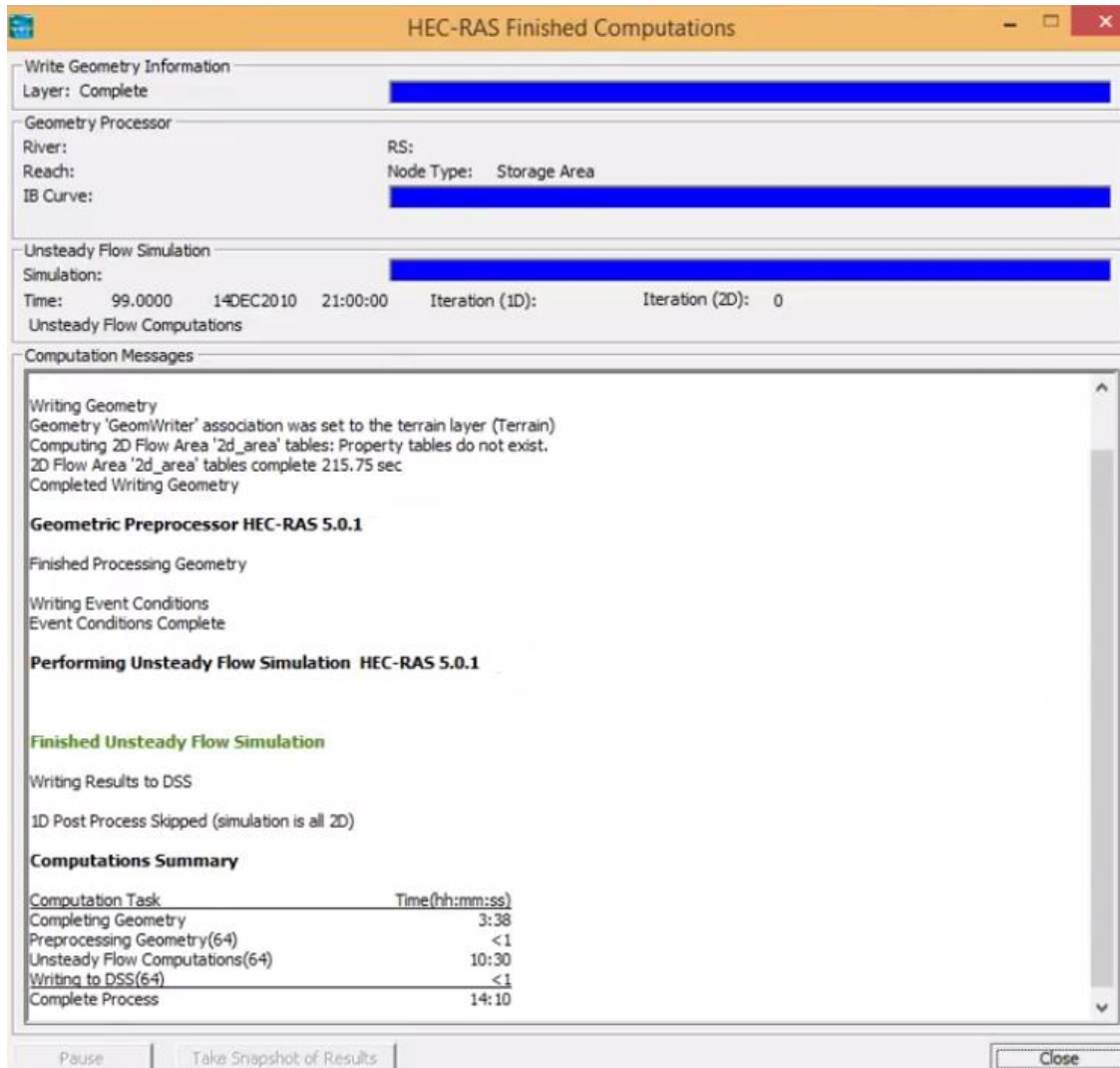


Fig 4.28 HEC RAS COMPUTATIONS

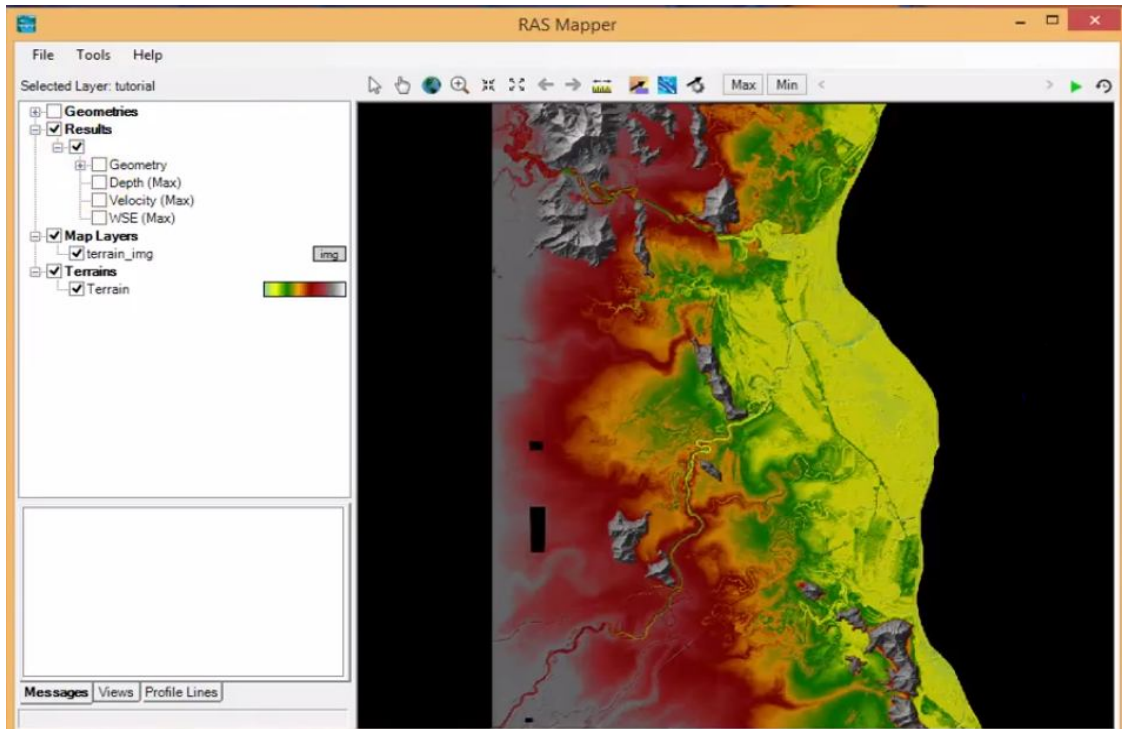


Fig 4.29 COMPUTATION RESULTS

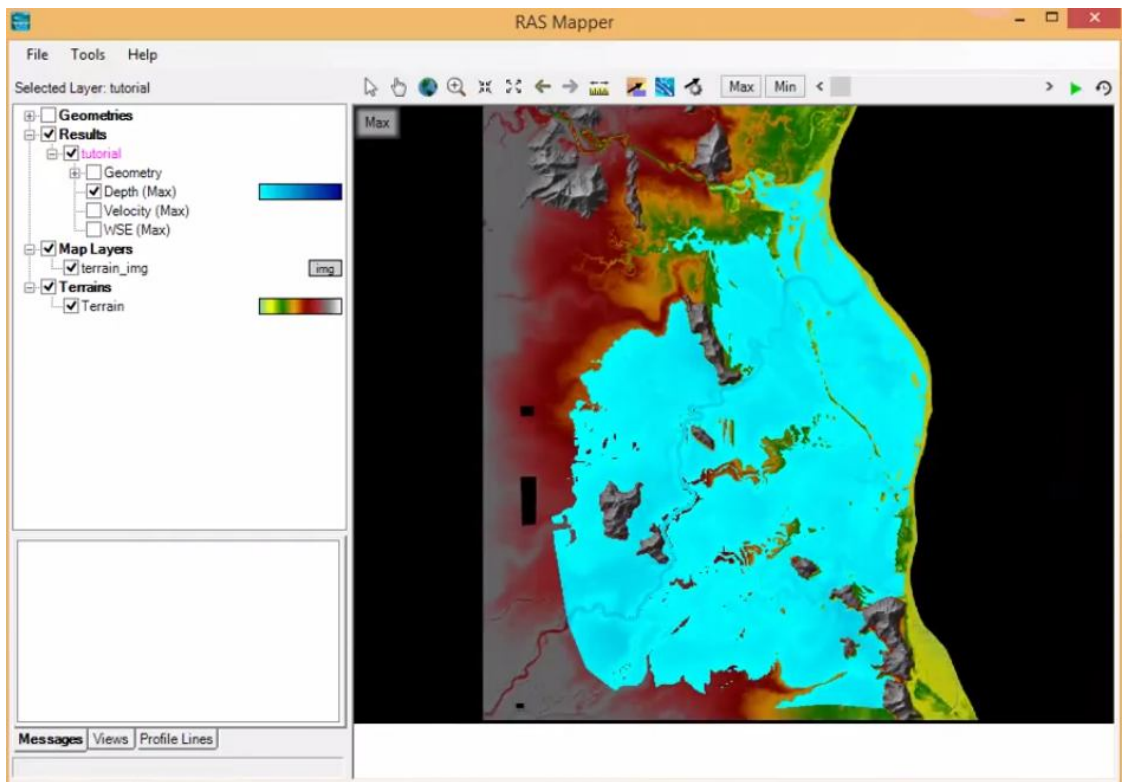


Fig 4.30 DEPTH OF INUNDATION MAP FOR 100 YEAR FLOOD

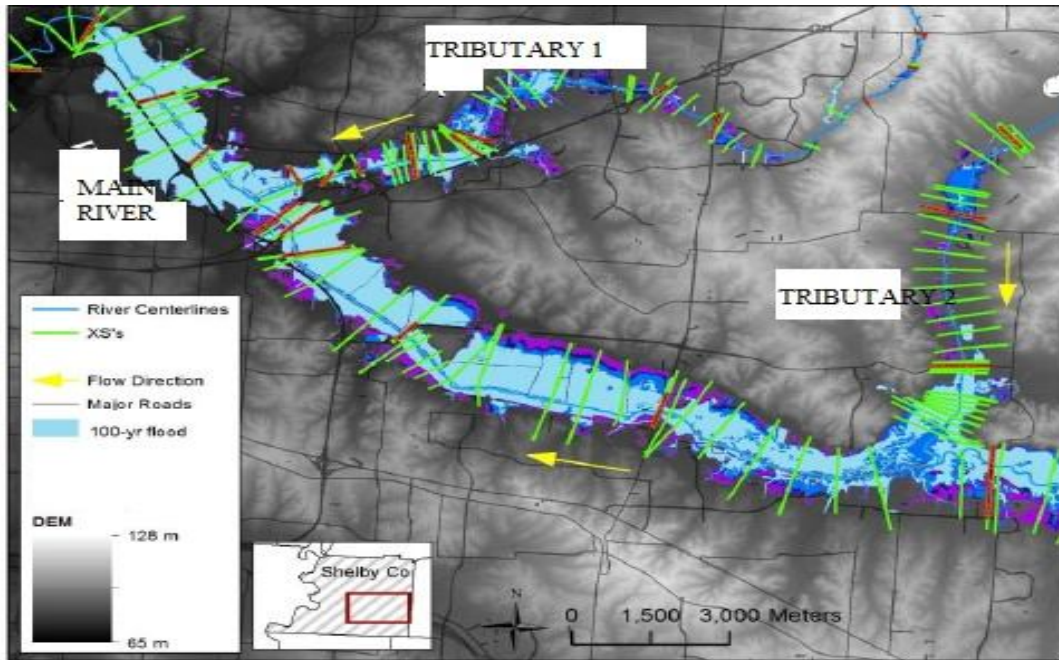


FIG 4.31 – FLOOD EXTENT WITH OVERLAY IMAGE OVER GROUND

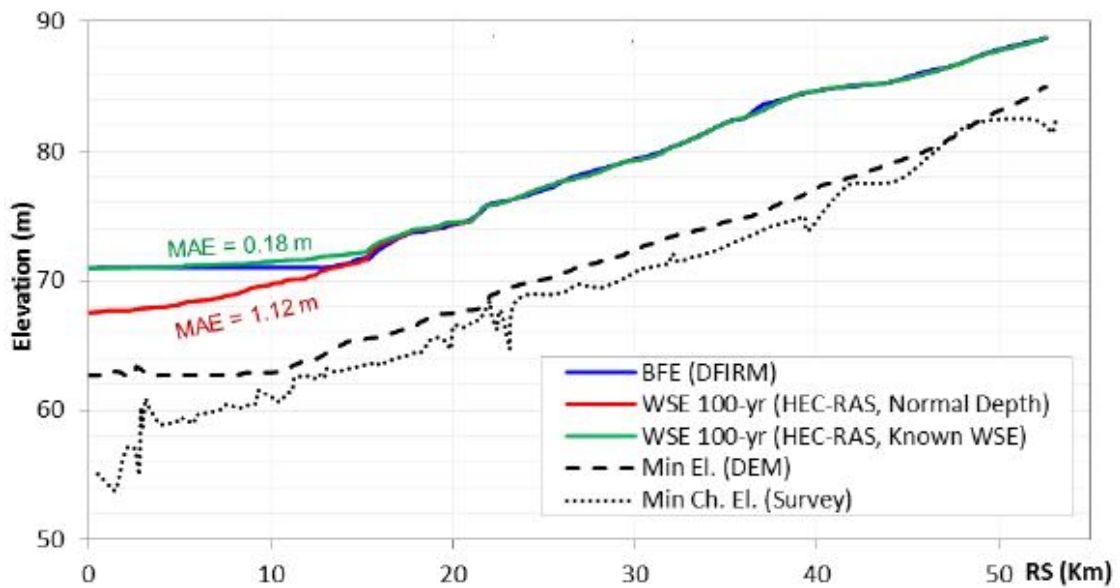


FIG 4.32 MAIN RIVER WATER DEPTH ALONG RIVER

RED LINE-WATER DEPTH FOR KNOWN NORMAL DEPTH FOR 100 YEAR RETURN PERIOD

GREEN LINE- KNOWN DEPTH ELEVATION FOR KNOWN WATER SURFACE ELEVATION

DOTTED LINE – SHOWS MINIMUM CHANNEL ELEVATION

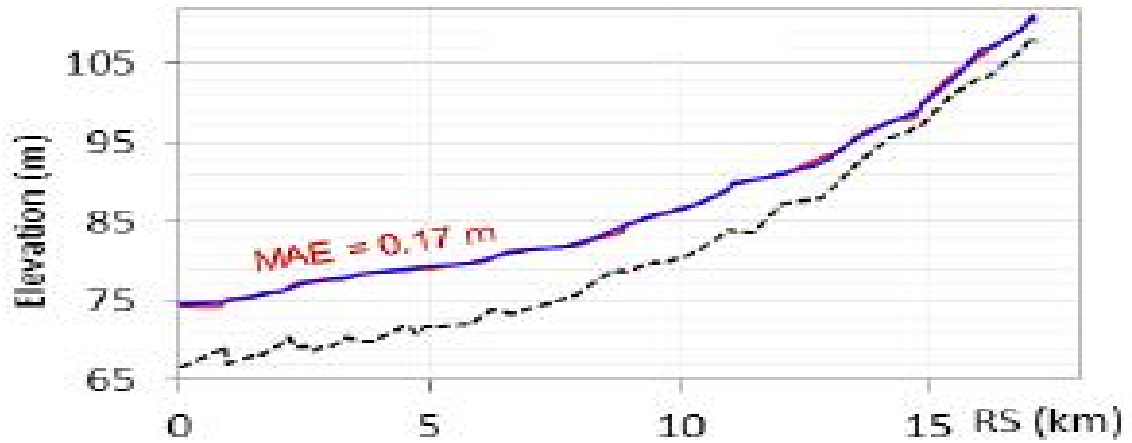


FIG 4.33 TRIBUTARY 1 WATER DEPTH

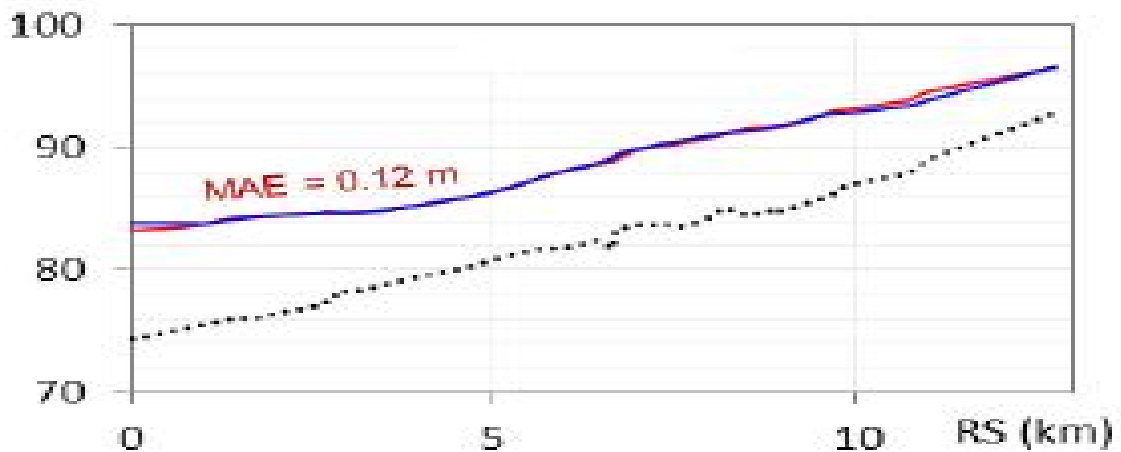


Fig 4.33 TRIBUTARY 2 WATER DEPTH ALONG THE RIVER

RED LINE –WATER DEPTH FOR KNOWN NORMAL DEPTH

BLUE LINE- WATER DEPTH FOR KNOWN WATER SURFACE ELEVATION

DOTTED LINE SHOWS –MIN CHANNEL ELEVATION



FIG 4.35 SHOWING MAIN RIVER FLOOD ON ITS BANKS

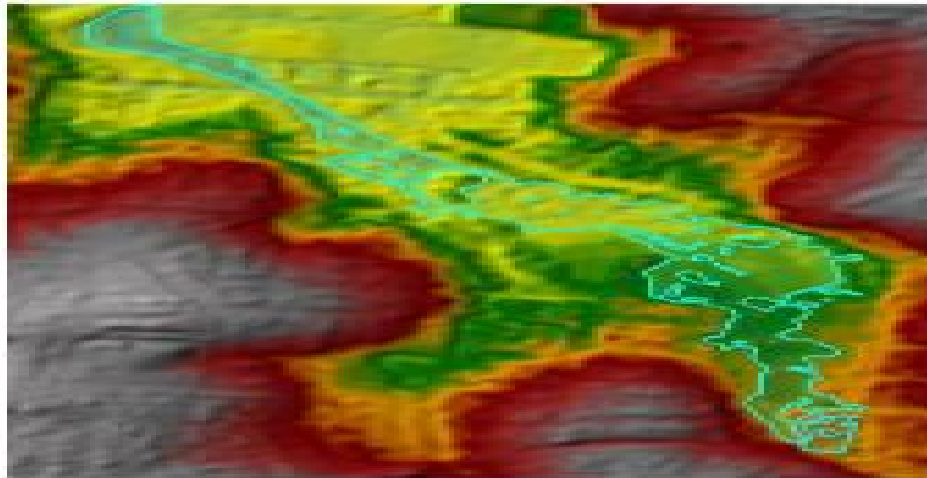


FIG 4.34 – SHOWING MAXIMUM DEPTH CONTOUR IN FLOODED AREA

4.36 FLOODING CHARACTERISTICS OF THE STUDY AREA

	For known normal depth	For known water surface elevation
FLOW VOLUME(1000M³)	1888338.20	1522613.16
FLOOD AREA(KM²)	128.20	115.89
INUNDATION DEPTH(m)	MAX.- 2.01 MIN-.35 MEAN- 1.05	MAX.- 1.85 MIN-.24 MEAN- .93
FLOW VELOCITY(M/S)	MAX-3.51 MIN-.01 MEAN-1.25	MAX-3.27 MIN-.01 MEAN-1.28

TABLE 6

CHAPTER 5

CONCLUSIONS

- When 100 year flood is evaluated max flow depth in the flood plain in the entire stretch that was studied is 201 cm and 185 cm respectively
- Max velocity found in the entire stretch 3.51 m/s and 3.27 m/s for two downstream condition.
- Area of inundation vary for main river stretch between 115-130 km²
- Max Distance of up to which inundation occurs over right bank is 600 m and for left bank is 1500 m.

When inundation areas are examined, it is seen that mostly agricultural areas and small parts of several little settlements and some major roads, are submerged. In study area, most of flood damage is originated from agricultural losses.